

# The detection of electromagnetic processes in the ionosphere caused by seismic activity

S.V.Koshevaya<sup>1</sup>, R. Pérez-Enríquez<sup>2</sup> and N. Ya. Kotsarenko<sup>2</sup>

<sup>1</sup> Instituto Nacional de Astrofísica, Óptica y Electrónica (INAOE), Puebla, Mexico.

<sup>2</sup> Depto. de Física Espacial, Instituto de Geofísica, UNAM, México, D. F., Mexico.

Received: February 22, 1996; accepted: June 19, 1996.

## RESUMEN

El propósito del presente artículo es el de presentar una nueva dirección en la investigación de terremotos por medio de procesos electromagnéticos en la ionosfera causados por actividad sísmica. Con este fin, se consideran las principales perturbaciones de la ionosfera causadas por terremotos, se analizan los mecanismos de conexión entre la litosfera y la ionosfera, y se discute la posibilidad de establecer una alarma para la ocurrencia de terremotos por medio de métodos espaciales y terrestres.

**PALABRAS CLAVE:** Ionosfera, actividad sísmica.

## ABSTRACT

The purpose of this article is to present a new direction in the investigation of earthquakes through the research of electromagnetic processes in the ionosphere caused by seismic activity. To accomplish this objective, the main perturbations of the ionosphere caused by earthquakes are considered, the connection mechanisms between the lithosphere and the ionosphere are analysed, and the possibility to establish a warning for the occurrence of earthquakes by means of space and ground methods is discussed.

**KEY WORDS:** Ionosphere, seismic activity.

## 1. INTRODUCTION

The prediction of earthquakes and volcanic eruptions remains largely an unsolved problem. Disastrous earthquakes, which occur from 100 to 200 times per year, are a hazard for every other inhabitant of our planet. Annually, thousands of people lose their lives under collapsed buildings, in fire and tsunamis caused by earthquakes.

## 2. PERTURBATION OF THE IONOSPHERE CAUSED BY EARTHQUAKES

A system of monitoring the occurrences of earthquakes from space must rely on the connection between the lithosphere, the ionosphere and the magnetosphere of the Earth. This connection may be established either from ground or from space.

There is at present factual material that shows evidence of a response in the ionosphere from seismic activity. Above the epicentre of a future earthquake, at altitudes from about 400 km to about 1000 km in the ionosphere, there appear macroscopic changes of the ionospheric parameters prior to the occurrence of the earthquake.

The presence in the ionosphere of precursors of earthquakes affords the possibility in principle of prediction by remote space-ground methods. The study of the influence of the seismic activity on the ionosphere began more than 30 years ago (see references). Some of the observations are the following.

- An anomaly in the absorption of cosmic ray emission after the 1960 earthquake in Chile lasted for 6 days (Warwick, 1963).

- Strong changes in the parameters of the ionosphere caused by the large earthquake in Alaska in 1964 (Davies and Barker, 1965).

- The discovery by Tarantsev and Birfeld (1973) that acoustic waves are involved in the connection between seismic activity and the ionosphere.

Since 1975, when the Soviet-French experiment "Arkad" was active in the ionosphere and the magnetosphere, processes caused by seismic activity were observed. The identification of these processes took place not at the time of the earthquake but hours or days before. Similar phenomena were observed on "Intercosmos 19", "Intercosmos 24", "OGO-6", "Nimbus", "GEOS 1", "GEOS 2", and other satellites. We propose five main types of ionospheric perturbations that accompany earthquakes.

(1). Variations of electric and magnetic fields.

The variations of electric  $\delta E$  and geomagnetic  $\delta H$  fields, as well as their frequency ranges are:

$$\delta E \approx 10^{-7} - 10^{-2} \text{ V/m}, f \approx 10^{-2} - 10 \text{ Hz},$$
$$\delta H \approx 10^{-4} - 10^{-1} \text{ nT}, f \approx 10^{-2} - 10 \text{ Hz}.$$

Variations of the electric field caused by earthquakes were first proposed by Chernyavskiy (1925). Variations of the magnetic field  $\delta H$  were observed in Kazan in 1880.

(2). Perturbations of the electromagnetic waves at extremely low (ELF) and very low frequencies (VLF) as follows.

(ELF)  $f \approx 10^{-2} - 10^2$  Hz,  $I_f$  maximum at  $f \approx 8$  Hz;

(VLF)  $f \approx 10^2 - 10^5$  Hz,  $I_f$  maximum at 10-15 KHz;

where  $I_f$  is the intensity of the electromagnetic perturbation.

(3). Perturbations of density and temperature of the ionosphere plasma in the F- and E -layers. In some cases the density may increase and in others it may decrease.

The observed values of the concentration in F-layer are:

$$\begin{aligned} \delta n / n &\approx 10^{-3} - 10^{-1}, \\ \delta T > 0, \delta T / T &\approx 10^{-2} - 10^{-1} \end{aligned}$$

where  $\delta n$  is the density change and  $T$  is the temperature.

(4). Increases of the intensity of luminescence of the ionosphere at the main spectral wavelengths of atomic oxygen  $\lambda = 5577 \text{ \AA}, 6300 \text{ \AA}$ .

Ionospheric luminescence before earthquakes is well known. It occurred above Rome in 373 BC. This phenomenon has also been reported above Tashkent and Spitak, Armenia during the 1966 and 1988 earthquakes, respectively.

(5). Flows of geoactive particles in the magnetosphere.

It turns out that the variations of the  $E, H$  fields, the low-frequency oscillations and the presence of flows of geoactive particles appear as a rule some hours before and up to the beginning of the earthquake, whereas the variations of the density, the temperature and the optical emission may appear one or two days before and until the occurrence of the earthquake.

Other phenomena have been observed, such as anomalies in the propagation of radio waves above the epicentre of earthquakes (Fuks and Shubova, 1994); perturbation of the sporadic Es-layer emission in the ionosphere; geochemical and biological processes in the seas, oceans, and others.

### 3. MAIN MECHANISMS OF CONNECTION BETWEEN LITHOSPHERE AND IONOSPHERE

At present no consistent theory has been developed about the connection between the lithosphere and the ionosphere. The mechanisms of interaction between these regions are not fully understood. However, it is clear that the main mechanisms of transfer of energy from the lithosphere to the ionosphere must be acoustic or electromagnetic.

The mechanics of the lithosphere-ionosphere connection involve the transformation of acoustic waves into magnetoacoustic and Alfvén waves (Figure 1). The oscillations of the Earth's surface excite acoustic and gravity waves in the atmosphere. After they spread upwards they are transformed into Alfvén and magnetoacoustic waves. They appear as electromagnetic oscillations of low fre-

quency. Consider the equations of the theory of elasticity for the lithosphere:

$$\rho \frac{\partial^2 \bar{U}}{\partial t^2} = \lambda \nabla \operatorname{div} \bar{U} + \mu \Delta \bar{U} + \bar{F} \quad (1)$$

where  $\bar{U}$  is the acoustic displacement,  $\rho$  is the density,  $\lambda$  and  $\mu$  are the coefficients of elasticity, and  $\bar{F}$  is the force of electroacoustic origin which is the source of excitation of acoustic waves. The equations of hydrodynamics for the atmosphere are:

$$\rho \left( \frac{\partial \bar{v}}{\partial t} + (\bar{v} \cdot \nabla) \bar{v} \right) = -\nabla P + \rho \bar{g}, \quad (3)$$

$$\frac{\partial \rho}{\partial t} + \operatorname{div} (\rho \bar{v}) = 0, \quad (4)$$

$$P = P(\rho, T), \quad (5)$$

where  $\bar{v}$  is the velocity,  $P$  is the pressure,  $\bar{g}$  is the acceleration of gravity and  $T$  is the temperature, and the equations of magnetohydrodynamics for the ionosphere are:

$$\rho \left( \frac{\partial \bar{v}}{\partial t} + (\bar{v} \cdot \nabla) \bar{v} \right) = -\nabla \left( p + \frac{H^2}{8\pi} \right) + \frac{(\bar{H} \cdot \nabla) \bar{H}}{4\pi} \rho \bar{g}, \quad (6)$$

$$\operatorname{rot} \bar{H} = \frac{4\pi}{c} \bar{j}, \bar{j} = \hat{\sigma} \cdot \bar{E}, \quad (5)$$

$$\operatorname{rot} \bar{E} = -\frac{1}{c} \frac{\partial \bar{H}}{\partial t}, \quad (7)$$

where  $\bar{E}, \bar{H}$  are electromagnetic fields, and  $\hat{\sigma}$  is the conductivity tensor of the ionosphere. It is necessary to introduce boundary conditions, which requires computer modelling.

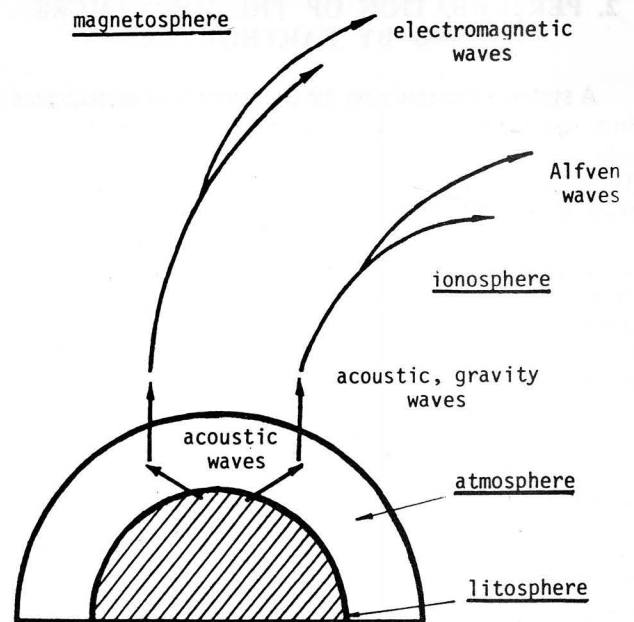


Fig. 1. Connection between the lithosphere and the ionosphere involving the transformation of acoustic waves into Alfvén waves, magnetoacoustic waves and others.

Another model was proposed by Schuman 1952, (Figure 2). The nonstationary electrical currents in the lithosphere are connected with the horizontal motion of the ground having a relatively high conductivity. According to Maxwell's equations, it becomes almost instantly apparent in the ionosphere. Suppose that there is a known current determined by geophysicists in a region of the lithosphere. Then there should exist corresponding electromagnetic fields in the ionosphere (Molchanov *et al.*, 1993). If the Earth's surface and the ionosphere constitute an electromagnetic resonator (Schuman's resonator), this current excites electromagnetic oscillations. Thus the maximum of low frequency electromagnetic oscillations found through satellite observations occurs approximately at a frequency  $f \approx 8\text{Hz}$ , which corresponds to the first characteristic frequency of Schuman's resonator (Schuman, 1952):

$$f \approx \frac{c}{2\pi R} \sqrt{n(n+1)}, \quad n=1,2,3 \quad (8)$$

where  $R$  is the radius of the Earth (Figure 2).

The maximum of VLF oscillations falls in the frequency range of  $f \approx 10 - 15 \text{ KHz}$ , corresponding to transverse oscillations of the resonator formed by the surface of the Earth and the ionosphere:

$$f \approx \frac{nc}{L}, \quad n=1,2,3 \quad (9)$$

where  $L$  is the altitude of the  $D$  - layer of the ionosphere. The spectrum of the oscillations is shown in Figure 2. The observations give the same frequencies (Figure 3).

The use of a space system for monitoring seismic activity presents some difficulties. It is necessary to take into account:

(a) changes in solar activity capable of generating similar

signals in the ionosphere. These are considered by using available solar geophysical data;

(b) anthropogenic factors (artificial explosions, launching of rockets and so on);

(c) the short time available for the treatment of satellite information.

Plasma experiments may be carried out in the ionosphere, above the area of seismic activity, during periods of months or years, but the effects of seismic activity on ionospheric parameters will be felt minutes or hours, at best, before the earthquake.

#### 4. WARNING FOR EARTHQUAKES

None of the satellites was designed for the investigation of processes in the ionosphere caused by seismic activity. The results that have been obtained in this field are fortuitous. The Ukrainian space project "Warning" whose scientific leader until 1995 was one of us (N. Y. K.) will be devoted to these investigations, and specifically the electromagnetic processes in the ionosphere caused by seismic activity (Kotsarenko *et al.*, 1995). This project is in charge of the Kiev National University (Department of Astronomy and Space Physics) and Pivdenne (Yuzhnoe) Design Bureau (Dnepropetrovsk). The participating countries are Russia, Austria, Hungary, Poland, Czech Republic, Romania, France, Germany, and Ukraine. The project includes rocket assembly, rocket and launching facilities; ground-based flight control with telemetry and communication links and stations; ground-based high-performance working in close connection with the spaceborne computer during data acquisition and processing; and a network of seismological stations and other ground instrumentation such a ionospheric and magnetic stations, optical facilities for the ground-based measurements and so on.

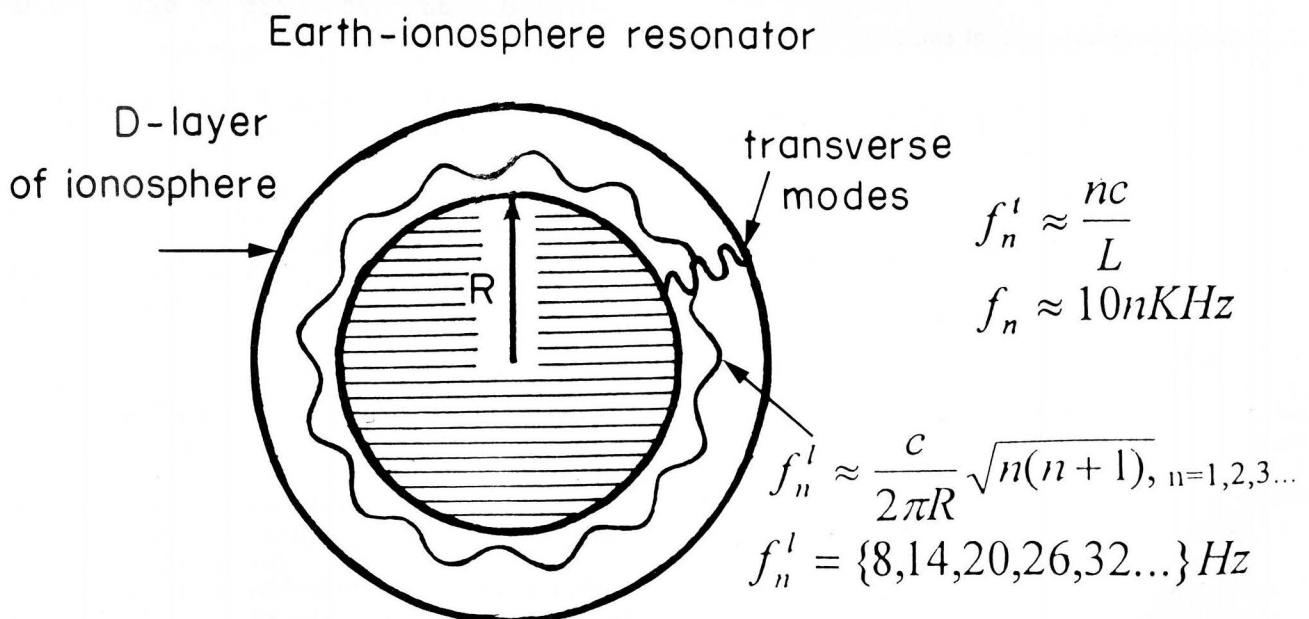


Fig. 2. Schuman resonator of the Earth-ionosphere cavity.

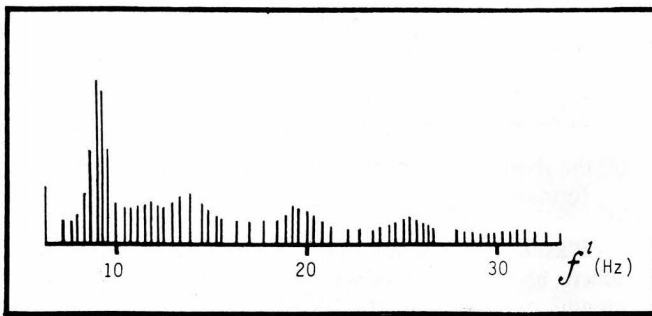


Fig. 3. The spectrum of Schuman oscillations obtained from ground observations.

The parameters of the space platform are shown below.

orbit	- circular, 600 km
declination	- 74 degrees
composition	- a satellite +2 sub-satellites
weight	- satellite-1500 kg; sub-satellite-100 kg
Satellite orientation accuracy	- better than 15 Deg
angular stabilisation velocity	- less than 0,05 Deg/s
period of active operation	- more than 1 year
rocket type	- Cyclone
tentative launching time	- 1998

The scientific payload of the spacecraft will contain the following equipment:

- Wave complex for measuring electric and magnetic fields;
- Instrumentation for measuring temperatures and concentrations of ionospheric plasma;
- Ionosonde radio-spectrometer for measuring electron concentration profiles below the F- layer maximum;
- System for measuring optical emissions of the ionosphere;
- Spectrometers of energetic particles (electrons, protons, ions);
- Two-frequency transmitter (150/400 MHz) for tomographic sounding of the ionosphere. The scheme of the space control system is shown in Figure 4.

The global character of the seismic problem calls for remote sensing methods of investigation. Only remote methods (mainly space borne) can ensure obtaining the necessary information from large regions with sufficient time resolution (Mc Farland et al., 1990; Alishouse et al., 1980; 1990; Barrett et al., 1981; Gasiewski et al., 1990; Neale et al., 1990; Petty and Kotsaros, 1990; Westwalter et al., 1989).

It is important to use appropriate ground control such as the mountain landmarks or the ocean coast. Radiometers, bolometers, and magnetometers on mountain tops will

provide a reliable surveillance of the environment. Sensing from space does not always provide the temporal and spatial resolution required. Furthermore, it is rather expensive and requires equipment calibration, which is impossible without the use of ground control. The radiometers, bolometers, and magnetometer remote-sensing systems have certain advantages when compared with other spaceborne methods:

- (1). The space remote sensing does not always provide good space-time resolution (Basharrimov et al., 1970; Hollinder et al., 1990).
- (2). The space remote sensing needs calibration by ground observations (Hollinder et al., 1990).
- (3). The spaceborne sensing does not always provide fast control since it depends on weather conditions (Goodbelret, 1990).

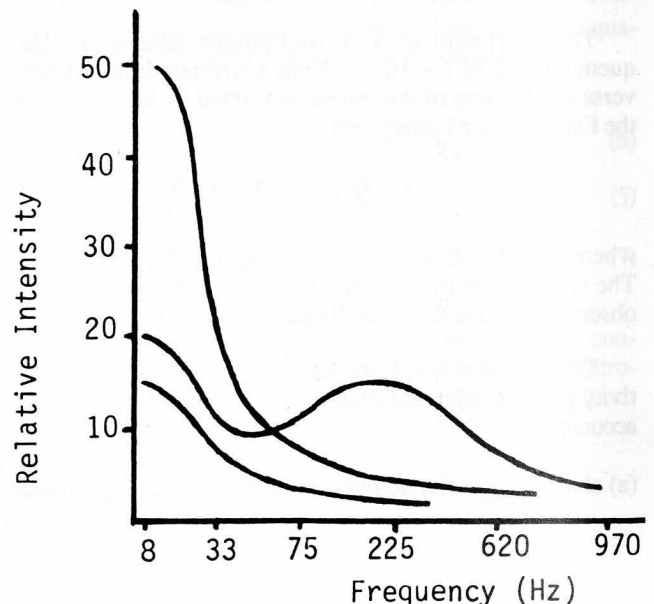


Fig. 4. The spectrum of Schuman oscillations obtained from satellite observations.

A Mexican project by Instituto Nacional de Astrofísica, Óptica y Electrónica-INAOE, together with Instituto de Geofísica-UNAM devoted to creation to a radiometer, bolometer, and magnetometer remote-sensing system and computer warning for earthquakes is being proposed as follows.

The radioelectronic system will be created on the basis of integrated low-noise receivers (Koshevaya et al., 1982) with antennas in different frequencies controlled by a personal computer. The data acquisition, the observation analysis, and the forecast of weather phenomena are based on the computer and spatial programs. The system will be based on an analogue system created at the Saturn Institute in Kiev (Chmil et al., 1995). This Institute has participated in design of new receivers for radiotelescopes and remote space communication, some of which are working

successfully on the Ratan radiotelescope and in the centre of remote space communication in Jevpatoria, Ukraine (Bakitko *et al.*, 1993).

This ground-based remote-sensing system has a number of advantages when compared with satellite sensing:

- (1) The use of identical elements gives the possibility to eliminate device errors, to increase reliability and precision of measurements and to simplify calibration.
- (2) The system is not as expensive as space systems, as it uses identical integrated elements with a small energy consumption.
- (3) The use of mountain landmarks with simultaneous meteorological stations and a net of radiometers, bolometers, and magnetometers enables us to increase the number of observations and to improve control of parameters of sensing.
- (4) The proposed system is ecologically safe, indestructible and non-polluting (Guiraud, 1979; Snider *et al.*, 1980).

The remote-sensing system is based on measuring the characteristics of radioheating radiation of the atmosphere and ocean in the microwave range and low frequency magnetic fields. The basic parameters are the radiobrightness temperature in the microwave range, the temperature contrasts and the value of magnetic field.

In the remote-sensing system, the use of new types of integrated elements (Chmil *et al.*, 1995; Bakitko *et al.*, 1993) is considered, with low-cost technology and a new method for measuring ocean parameters and vertical profiles of wind velocity.

#### 4. CONCLUSION

Measuring the radiobrightness temperature and magnetic fields may provide warning capabilities for earthquakes. Microwave radiation at all frequencies generated under the epicentre causes the radiobrightness temperatures of the ionosphere and atmosphere as well as the surface at the epicentre to change very rapidly. This provides the possibility together with seismic and meteorological stations, to monitor a region for future volcanic and seismic phenomena.

We conclude that, in principle, the accumulated experience through observations on the ground and in space may raise the possibility of creating a global space-ground system for detection and prediction of earthquakes.

#### BIBLIOGRAPHY

- ALISHOUSE, J. C., J. B. SNIDER, E. D. R. WESTWATER, C. T. SWIFT, C.S. RUF, S.A. SNYDER, J. VONGSATHORN and R.R. FERRARO, 1980. Determination of Ocean Total Precipitation Water from the SSM/I. *IEEE Trans. on Geoscience and Remote Sensing*, 28, 817-822.
- ALISHOUSE, J.C., R. FERRARO and J.V. FIORE, 1990. Influence of Oceanic Rainfall Properties from the Nimbus-7 SMMR. *J. Appl. Meteor.*, 29, 551-559.
- BAKITKO, R.V.M., M.B. VASILYEV, A.S. WINITZKI, *et al.*, 1993. Radiosystems of Remote Space Communication. Moscow. Radio and Communication Publishing House, 1993, 328p. (in Russian).
- BARRETT, E. C. and D. W. MARTIN, 1981, The Use of Satellite Data in Rainfall Monitoring. Academic Press. 340p.
- BASHARRIMOV, A. E., A. S. GURVICH, L. T. TUCHKOV and K. S. SHIFRIN, 1970. The Terrestrial Thermal Radio-Emission Field. *Atmos. Ocean. Phys.*, 6, 210-218.
- CHERNYAVSKY, E., 1925. Electric Storms. Proc. of CASU, p.157. (in Russian).
- CHMIL, V. M., K. S. SUNDUCHKOV and L. S. NASARENKO, 1995. The Border Multichannel Radiometrical Complex for the Distant Sounding of the Earth from the Space RMC/KII. Proc. of the 5th International Symposium on Recent Advances in Microwave Technology (ISRAMT'95), Kiev, p.141-146.
- DAVIS, K. and D. BARKER, 1965. Ionospheric Effects Observed Around the Time of the Alaska Earthquake of March, 1964. *J. Geophys. Res.*, 70, 2551-2553.
- FUKS, I. M. and R. S. SHUBOVA, 1994. ELF-signal anomalies as a response of the lower ionosphere to conductivity change in the atmosphere. *Geomagnetism and Aeronomy*, N2, 130.
- GASIEWSKI, A. J., J. W. BARRETT, P. G. BONANNI and D. H. STAELIN, 1990. Aircraft-Based Radiometric Imaging of Tropospheric Temperature and Precipitation Using the 118.75 GHz Oxygen Resonance. *J. Appl. Meteor.*, 29, 620-632.
- GOODBELRET, M. A., 1990. Ocean Surface Wind Speed Measurements of the Special Sensor Microwave/Imager (SSM/I). *IEEE Trans. Geoscience, Rem. Sens.*, 28, 823-828.
- GUIRAUD, F. O., 1979. A Dual-Channel Microwave Radiometer for Measurement of Precipitable Water Vapor and Liquid, *IEEE Trans. Geoscience Electr.*, GE-17, 551-554.
- HOGG, D. C., 1980. Design of a Ground-Based Remote Sensing System Using Radio Wavelengths to Profile Lower Atmospheric Winds, Temperature and Humidity. *Remote Sensing of Atmosphere and Oceans*, Academic Press Inc., 313-364.
- ALISHOUSE, J. C., J. B. SNIDER, E. D. R. WESTWATER, C. T. SWIFT, C.S. RUF, S.A. SNYDER, J. VONGSATHORN and R.R. FERRARO, 1980. Determination of Ocean Total Precipitation Water from

- HOLLINDER, J. P., J. L. PEIRCE and G. A. POE, 1990. SSM/I Instrument Evaluation. *IEEE Trans. Geoscience Rem. Sens.* 28, 781-789.
- KOSHEVAYA, S. V., B. N. EMELYANENKOV and L. G. GASSANOV, 1982. Integrated Circuits for the MM Range. *Radioelectr. Com. Systems*, 25, 14-31.
- KOTSARENKO, N. Ya., V. E. KOREPANOV and V. N. IVCHENKO, 1995. The Investigation of the Ionosphere Precursors of Earthquakes (Experiment "Warning"). *Kosmichnaya nauka i tehnologiya*, 1, 93-96. (in Russian).
- MOLCHANOV, O. A., O. A. MAZHAEVA, A. N. GOLIAVIN and M. HAYAKAWA, 1993. Observation by the Intercosmos-24 Satellite of ELF-VLF Electromagnetic Emission Associated with Earthquakes. *Ann. Geophysics* 11, 431-440.
- McFARLAND, M. J., R. L. MILLER and C. M. U. NEALE, 1990. Land-Surface Temperature Derived from the SSM/I Passive Microwave Brightness Temperatures. *IEEE Trans. Geoscience Rem. Sens.*, 28, 939-945.
- NEALE, C.M.U., M.J. McFARLAND and KAI CHANG, 1990. Land-Surface-Type Classification Using Microwave Brightness Temperatures from the Special Sensor Microwave/Imager. *IEEE Trans. Geoscience Rem. Sens.*, 28, 829-838.
- PETTY, G.W. and K.B. KATSAROS, 1990. Precipitation Observed over the South China Sea by the Nimbus-77 Scanning Multichannel Microwave Radiometer during Winter MONEX. *J. Appl. Meteor.* 29, 273-287.
- SCHUMAN, W.O., 1952. Über die strahlungslosen Eigenschwingungen einer leitenden Kugel, die von einer Luftschicht und einer Ionosphärenhülle umgeben ist. *Z. Naturforsch.* 7a, 149-154.
- SNIDER, J.B., F.O. GUIRAND and D.C. HOGG, 1980. Comparison of Cloud Liquid Content Measured by Two Independent Ground Based Systems. *J. Appl. Meteor.*, 19, 577-579.
- TARANTSEV, A. and Ya BIRFIELD, 1973. Discovery: Influence of the Seismic Activity on Ionosphere by Acoustic Waves. *Short description of discoveries*, 128, 157. (in Russian).
- WARWICK, J. W., 1963. Radioastronomical Techniques for the Studies of the Atmosphere. (Ed. by J. Aarons). North Holland, Amsterdam p. 400.
- WESTWALTER, E. D., 1989. Combined Ground-and Satellite Based Radiometric Remote Sensing. In: "RSRM" 87: Advances in Remote Sensing Retrieval Methods, DEEPAK Publishing, 215-228.
- 
- S.V. Koshevaya<sup>1</sup>, R. Pérez-Enríquez<sup>2</sup> and N. Ya. Kotsarenko<sup>2</sup>
- <sup>1</sup> Instituto Nacional de Astrofísica, Óptica y Electrónica (INAOE), P.O.BOX 518&216, 7200 Puebla, México.  
e-mail: svetlana@inaoep.mx
- <sup>2</sup> Depto. de Física Espacial, Instituto de Geofísica, UNAM, Coyoacán, 04510 México, D. F., México.