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LOVE WAVE IN POROUS LAYER UNDER INITIAL STRESS OVER HETEROGENEOUS ELASTIC HALF-SPACE UNDER GRAVITY AND INITIAL STRESS

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RESUMEN

En este trabajo se estudió el efecto de las tensiones iniciales y la gravedad sobre la propagación de las ondas de Love, lo anterior en la superficie de una capa porosa sobre un semiespacio heterogéneo. Se consideraron dos tipos de límite en superficies libres: (a) límite rígido y (b) límite libre de tracción. La propagación de las ondas de Love ha sido investigada bajo supuestos medios, tanto en los casos de frontera como en un estudio comparativo de dos casos. En ambos, se obtuvieron las ecuaciones de dispersión y las velocidades de fase. Se presentan los cálculos numéricos de forma gráfica. Este estudio de las ondas de Love en el medio supuesto revela que la presencia de tensión inicial en el medio espacio y la ausencia de tensión inicial en la capa, el desplazamiento de la velocidad de fase en un límite rígido es mayor que el límite libre de tracción.

PALABRAS CLAVE: Ondas de Love, propagación, tensión inicial, heterogéneo, poroso y gravedad.

ABSTRACT

In the present paper, the effect of initial stresses and gravity on the propagation of Love waves have been studied in porous layer surface over a heterogeneous half-space. We have considered two types of boundary on free surfaces: (a) rigid boundary and (b) traction free boundary. The propagation of Love waves has been investigated under assumed media in both the cases of boundary and discusses a comparative study of two cases. The dispersion equations and phase velocities have been obtained in both cases. The numerical calculations have been done and presented graphically. This study of Love waves in the assumed medium reveals that the presence of initial stress in the half-space and absence of initial stress in the layer, the displacement of phase velocity in a rigid boundary is more than the traction free boundary.

KEY WORDS. Love wave, propagation, initial stress, heterogeneous, porous and gravity.

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INTRODUCTION

The stress is generally developed in media due to natural phenomena. Since the earth is an elastic solid medium under high initial stresses. These stresses play a significant role to propagate an elastic wave. Chattopadhyay et al. (1978) discussed the propagation of Love type waves in an initially stressed crustal layer having an irregular interface. Chakraborty et al. (1981) shows that Love waves propagate in dissipative media under gravity. Chakraborty et al. (1983) discussed the effect of initial stress and irregularity on the propagation of SH-waves. Dey et al. (1996) discussed the propagation of Love waves in heterogeneous crust over a heterogeneous mantle. Abd-Alla et al. (1999) have studied the propagation of Love waves in a non-homogeneous orthotropic elastic layer under initial stress overlying semi-infinite medium. Dey et al. (2004) also studied the propagation of Love waves in an elastic layer with void pores. Sharma (2004) established a mathematical expression about wave propagation in a general anisotropic poroelastic medium with anisotropic permeability phase velocity and attenuation. Kalyani et al. (2008) have made finite-difference modeling of seismic wave propagation in monoclinic media.

Many researchers in the field of elastic wave propagation in layered medium bounded by different forms of irregular boundaries have been studied in several research papers. Such as Anjana and Samal (2010) shows that Love waves propagate in a fluid-saturated porous layer under a rigid boundary and lying over elastic half-space under gravity. Chattaraj and Samal (2013) investigated the propagation of Love waves in the fiber-reinforced layer over a gravitating porous half-space. Gupta et al. (2013a) discussed the propagation of Love waves in a non-homogeneous substratum over an initially stressed heterogeneous half-space. Gupta et al. (2013b) were introduced the possibility of Love wave propagation in a porous layer under the effect of linearly varying directional rigidities. SH-type waves dispersion in an isotropic medium sandwiched between an initially stressed orthotropic and heterogeneous semi-infinite media were studied by Kundu et al. (2013). Manna et al. (2013) show the Love wave propagation in a piezoelectric layer overlying in an inhomogeneous elastic half-space. Bacigalupo and Gambarotta (2014) discussed second-gradient homogenized model for wave propagation in heterogeneous periodic media. Propagation of Love wave in fiber-reinforced medium lying over an initially stressed orthotropic half-space was formulated by Kundu et al. (2014).

In this paper, we have studied the problem of propagation of Love waves in a porous layer over a heterogeneous elastic half-space under gravitating half-space for a rigid boundary as well as traction free boundary in the upper layer. Both the layer is considered under the effect of initial stress. The dispersion relations have been derived for rigid boundary as well as traction free boundary. The initial stresses play a vital role on the propagation of Love waves in the assumed medium. Gravity and heterogeneity in the half-space play a notable effect on the propagation of Love waves in the medium and half-space. The influences of porosity, initial stresses, and gravitational parameters have discussed graphically.

FORMULATION OF THE PROBLEM

We consider a model consisting of the water-saturated anisotropic poroelastic layer under initial stress of finite thickness laying (Figure 1 under traction free boundary and Figure 2 under rigid boundary) over a gravitating heterogeneous elastic half-space under initial stress. Considering the origin of the coordinate system at the interface of the crust and mantle, z-axis downloads positively. The following variation has been taken.

For the half-space $\mu = \mu_2(1 + a_1z), \rho = \rho_2$

where, μ and ρ is the rigidity and density of the half space, a be the constant having inverse of length.

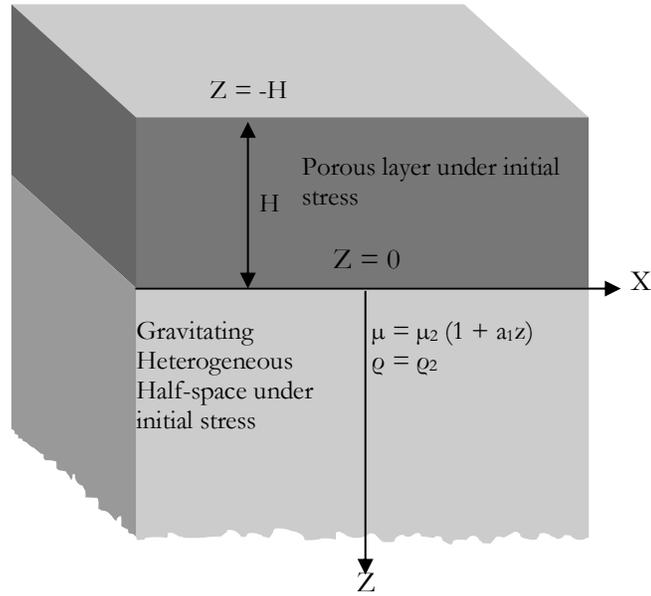


Figure 1: Geometry of the problem

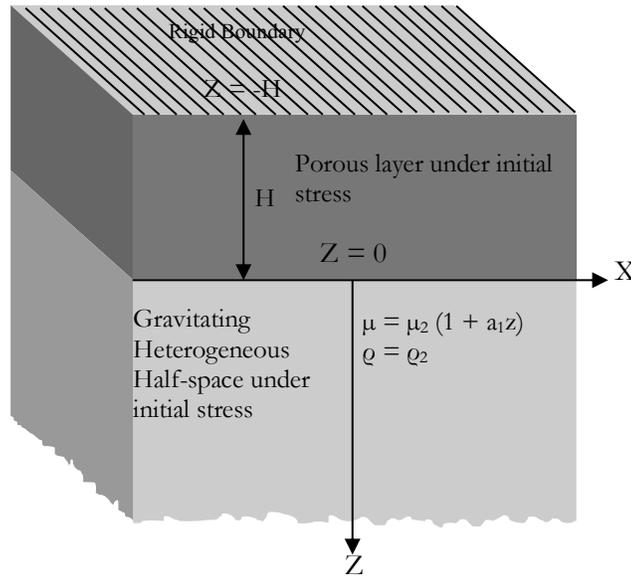


Figure 2: Geometry of the problem

SOLUTION OF THE POROUS LAYER

For the fluid-saturated anisotropic porous layer under initial stress P_1 in the absence of body forces, the equation of motion can be written as (Biot, 1965).

$$\begin{aligned} \frac{\partial s'_{21}}{\partial x} + \frac{\partial s'_{22}}{\partial y} + \frac{\partial s'_{23}}{\partial z} - P_1 \frac{\partial \omega_z}{\partial x} &= \frac{\partial^2}{\partial t^2} (\rho_{11} v'_y + \rho_{12} V_y) \\ \frac{\partial S}{\partial y} &= \frac{\partial^2}{\partial t^2} (\rho_{12} v'_y + \rho_{22} V_y) \\ s'_{23} &= 2L e_{yz}, \quad s'_{12} = 2N e_{xy}, \quad s'_{22} = (A - 2N) e_{xx} + A e_{yy} + F e_{zz} + Q \varepsilon \end{aligned} \tag{1}$$

where, s'_{ij} are the components of stress tensor in the solid, $S (= -fp)$ is the reduced pressure of the fluid, p is the pressure in the fluid, and f is the porosity of the porous layer, (u'_x, v'_y, w'_z) are the components of the displacement vector of the solid and (U_x, V_y, W_z) are those of fluid. L and N are represent the shear moduli of the anisotropic layer in the x - and z - direction respectively, whereas A and F are elastic constants for the medium. The positive quantity Q is the measure of coupling between the changes in the volume of solid and liquid.

Since the Love waves propagating along the x -direction, having the displacement of particles along the y -direction, we have

$$\begin{aligned} u'_x &= 0, \quad v'_y = v(x, z, t), \quad w'_z = 0 \\ U_x &= 0, \quad V_y = V(x, z, t), \quad W_z = 0 \end{aligned}$$

This displacement will produce e_{yz} and e_{xy} strain components and others are vanishing.

The dynamic components $\rho_{11}, \rho_{12}, \rho_{22}$ take into account the inertia effects of the moving fluid and are related to the densities of the solid ρ_s , fluid ρ_f and the layer ρ' by the equations

$$\rho_{11} + \rho_{22} = (1 - f) \rho_s, \quad \rho_{12} + \rho_{22} = f \rho_f$$

So that the mass density of the aggregate is

$$\rho' = \rho_{11} + 2\rho_{12} + \rho_{22} = \rho_s + f(\rho_f - \rho_s)$$

Also

$$\rho_{11} > 0, \quad \rho_{12} \leq 0, \quad \rho_{22} > 0, \quad \rho_{11} \rho_{22} - \rho_{12}^2 > 0.$$

The Love wave equation takes the form

$$\begin{aligned} \left(N - \frac{P_1}{2}\right) \frac{\partial^2 v}{\partial x^2} + L \frac{\partial^2 v}{\partial z^2} &= \frac{\partial^2}{\partial t^2} [\rho_{11} v + \rho_{12} V] \quad \text{and} \\ \frac{\partial^2}{\partial t^2} [\rho_{12} v + \rho_{22} V] &= 0 \end{aligned} \tag{2}$$

Hence, we have from (2)

$$\left(N - \frac{P_1}{2}\right) \frac{\partial^2 v}{\partial x^2} + L \frac{\partial^2 v}{\partial z^2} = d_1 \frac{\partial^2 v}{\partial t^2} \quad (3)$$

where $d_1 = \rho_{11} - \frac{\rho_{12}^2}{\rho_{22}}$, $V = (d_2 - \rho_{12}v) / \rho_{22}$, $d_2 = [\rho_{12}v + \rho_{22}V]$

The shear wave velocity in the porous layer along the x-direction can be expressed as

$$\beta' = \sqrt{\frac{\left(N - \frac{P_1}{2}\right)}{d_1}} = \beta_1 \sqrt{\left(1 - \frac{\xi_1}{d}\right)}$$

where $d = \gamma_{11} - \left(\frac{\gamma_{12}^2}{\gamma_{22}}\right)$, $\beta_1 = \sqrt{\frac{N}{\rho'}}$, is the velocity of shear wave in the corresponding initial stress free non-porous, anisotropic, elastic medium along x-direction.

$\xi_1 = \frac{P_1}{2N}$, is the non-dimensional parameter due to the initial stress P_1 and

$\gamma_{11} = \frac{\rho_{11}}{\rho}$, $\gamma_{12} = \frac{\rho_{12}}{\rho}$ and $\gamma_{22} = \frac{\rho_{22}}{\rho}$, are the non-dimensional parameter for the material of the porous layer as obtained by Biot (1965)

We consider

$$v(x, z, t) = f(z)e^{i(kx - \omega t)} \quad (4)$$

where, k is the wave number and ω is the angular frequency.

Now, we get

$$\frac{d^2 f(z)}{dz^2} + q_1^2 f(z) = 0,$$

where $q_1^2 = \frac{k^2}{L} \left(c^2 d - \left(N - \frac{P_1}{2} \right) \right) = k^2 \left\{ \alpha d \left(\frac{c^2}{\beta_1^2} - \frac{1 - \xi_1}{d} \right) \right\}$, $\alpha = \frac{N}{L}$, $\omega = kc$

(5)

The solution of equation (5)

$$f(z) = Ae^{iq_1 z} + Be^{-iq_1 z}$$

where A and B are constants (6)

Hence, finally we get

$$v_1(x, z, t) = (Ae^{iq_1 z} + Be^{-iq_1 z})e^{i(kx - \omega t)} \quad (7)$$

THE DYNAMIC EQUATION OF THE MOTION IN THE HALF-SPACE

The dynamic equation of motion in the half space may be written as Biot (1965)

$$\frac{\partial S_{12}}{\partial x} + \frac{\partial S_{22}}{\partial y} + \frac{\partial S_{23}}{\partial z} - \rho g \omega_{23} + \rho g z \frac{\partial \omega_{12}}{\partial x} - \rho g z \frac{\partial \omega_{23}}{\partial z} - \frac{P_2}{2} \frac{\partial v}{\partial x^2} = \rho \frac{\partial^2 v_2}{\partial t^2} \tag{8}$$

where, $v_2(x, z, t)$ is the displacement along y direction, ρ is the density, S_{ij} are incremental stress and ω_{ij} are rotational components in the half space and g is the acceleration due to gravity. P_2 , is the initial stress.

The components of body forces are $X = 0$, $Y = 0$, $Z = g$.

The stress-strain relations are

$$S_{12} = 2\mu e_{xy}, S_{23} = 2\mu e_{yz}, e_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$$

$$e_{xy} = \frac{1}{2} \left(\frac{\partial v_1}{\partial y} + \frac{\partial v_2}{\partial x} \right), e_{yz} = \frac{1}{2} \left(\frac{\partial v_2}{\partial z} + \frac{\partial v_3}{\partial y} \right)$$

In this problem $\frac{\partial}{\partial y} = 0$ and $\mu = \mu_2(1 + a_1 z)$, $\rho = \rho_2$, the stress strain relation becomes

$$e_{xy} = \frac{1}{2} \frac{\partial v_2}{\partial x}, e_{yz} = \frac{1}{2} \frac{\partial v_2}{\partial z}$$

$$S_{12} = \mu_2(1 + a_1 z) \frac{\partial v_2}{\partial x}, S_{23} = \mu_2(1 + a_1 z) \frac{\partial v_2}{\partial z}$$

(9)

where, μ is the modulus of rigidity, a_1 is a variation parameter of a rigidity having dimension inverse of length.

Equation (8) using the above relations takes the form

$$\left[\mu_2(1 + a_1 z) - \frac{1}{2} \rho_2 g z - \frac{P_2}{2} \right] \frac{\partial^2 v_2}{\partial x^2} + \left[\mu_2(1 + a_1 z) - \frac{1}{2} \rho_2 g z \right] \frac{\partial^2 v_2}{\partial z^2} + \left(a_1 \mu_2 - \frac{1}{2} \rho_2 g \right) \frac{\partial v_2}{\partial z} = \rho_2 \frac{\partial^2 v_2}{\partial t^2} \tag{10}$$

The solution of equation (10) may be taken as

$$v_2 = \phi(z) e^{i(kx - \omega t)} \tag{11}$$

Equation (10) is now

$$\left[\mu_2(1 + a_1 z) - \frac{1}{2} \rho_2 g z - \frac{P_2}{2} \right] [-k^2 \phi] + \left[\mu_2(1 + a_1 z) - \frac{1}{2} \rho_2 g z \right] \phi'' + \left(a_1 \mu_2 - \frac{1}{2} \rho_2 g \right) \phi' + \rho_2 c^2 k^2 \phi = 0$$

$$\phi'' + k \left\{ \frac{\frac{a_1}{k} - \frac{G}{2}}{1 + k \left(\frac{a_1}{k} - \frac{G}{2} \right) z} \right\} \phi' + k^2 \left[\frac{\frac{c^2}{2} - \frac{P_2}{2\mu_2}}{1 + k \left(\frac{a_1}{k} - \frac{G}{2} \right) z} - 1 \right] \phi = 0 \tag{12}$$

where $c_2^2 = \frac{\mu_2}{\rho_2}$, $\omega = kc$ and $G = \frac{\rho_2 g}{\mu_2 k}$, is the Biot's gravity parameter, $\phi'' = \frac{\partial^2 \phi}{\partial z^2}$, $\phi' = \frac{\partial \phi}{\partial z}$.

Now we substitute, $\varphi(z) = \frac{\phi(z)}{\left\{1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z\right\}^{\frac{1}{2}}}$ in equation (12), we have

$$\varphi'' + \left[\frac{k^2 \left(\frac{a_1}{k} - \frac{G}{2}\right)^2}{4 \left\{1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z\right\}^2} + k^2 \left(\frac{\frac{c^2}{c_2^2} - \frac{P_2}{2\mu_2}}{1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z} - 1 \right) \right] \varphi = 0 \quad (13)$$

Using dimensionless quantities, $\eta = -\frac{2\left[1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z\right]}{\left(\frac{a_1}{k}-\frac{G}{2}\right)}$

Equation (13) becomes

$$\frac{\partial^2 \varphi(\eta)}{\partial \eta^2} + \left[-\frac{1}{4} + \frac{R}{\eta} + \frac{1}{4\eta^2} \right] \varphi(\eta) = 0 \quad (14)$$

where $R = -\frac{\frac{c^2}{c_2^2} + \frac{P_2}{2\mu_2}}{2\left(\frac{a_1}{k}-\frac{G}{2}\right)}$

which is standard Whittaker's equation and solution is

$$\varphi(\eta) = DW_{R,0}(\eta) + EW_{-R,0}(-\eta) \quad (15a)$$

where, D and E are arbitrary constants and $W_{R,0}(\eta)$, and $W_{-R,0}(-\eta)$ are the Whittaker functions.

As the solution should vanish at $z \rightarrow \infty$, i.e. for $\eta \rightarrow -\infty$, we may take the solution as

$$\varphi(\eta) = W_{-R,0}(-\eta) \quad (15b)$$

Hence

$$v_2 = \frac{EW_{-R,0} \left\{ \frac{2\left[1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z\right]}{\left(\frac{a_1}{k}-\frac{G}{2}\right)} \right\}}{\left\{1+k\left(\frac{a_1}{k}-\frac{G}{2}\right)z\right\}^{\frac{1}{2}}} e^{i(kx-\omega t)} \quad (16)$$

BOUNDARY CONDITIONS AND DISPERSION RELATION

We consider two cases

Case A: If the surface of the upper layer is rigid boundary, then

- i. $v_1 = 0$ at $z = -H$
- ii. $(\Delta f_y)_1 = (\Delta f_y)_2$ at $z = 0$
- iii. $v_1 = v_2$ at $z = 0$

Case B: If the surface of the upper layer is traction free, then

- i. $(\Delta f_y)_1 = 0$ at $z = -H$
- ii. $(\Delta f_y)_1 = (\Delta f_y)_2$ at $z = 0$
- iii. $v_1 = v_2$ at $z = 0$

For case A using the boundary condition (i), we have

$$Ae^{-iq_1H} + Be^{iq_1H} = 0 \text{ at } z = -H \tag{17}$$

For case B using the boundary condition (i), we have

$$Ae^{-iq_1H} - Be^{iq_1H} = 0 \text{ at } z = -H \tag{18}$$

For case A and case B using the boundary condition (ii), we have

$$L[iq_1A - iq_1B] = \mu_2 \left. \frac{d}{dz} \left[\frac{EW_{-R,0} \left\{ \frac{2 \left[1 + k \left(\frac{a_1}{k} - \frac{G}{2} \right) z \right] \right\} e^{i(kx - \omega t)}}{\left(\frac{a_1}{k} - \frac{G}{2} \right)} \right]}{\left\{ 1 + k \left(\frac{a_1}{k} - \frac{G}{2} \right) z \right\}^{\frac{1}{2}}} \right] \right|_{z=0} \tag{19}$$

For case A and case B using the boundary condition (iii), we have

$$A + B - E[W_{-R,0}(-\eta)]_{z=0} = 0 \tag{20}$$

Case I -In rigid boundary, eliminating A, B and E from equation (17), (19) and (20), we have

Expanding the determinant, we get

$$\cot\left(\sqrt{\alpha d\left(\frac{c^2}{\beta_1^2}-\frac{1-\xi_1}{d}\right)}kH\right) = \frac{\sqrt{\alpha d\left(\frac{c^2}{\beta_1^2}-\frac{1-\xi_1}{d}\right)}\left[1-\frac{1}{8}\left\{\frac{-\left(\frac{c^2}{c_2^2}+\xi_2\right)}{\left(\frac{a_1-G}{k}-\frac{G}{2}\right)}+1\right\}^2\left(\frac{a_1-G}{k}-\frac{G}{2}\right)\right]}{\mu_2\left[\left\{1-\frac{1}{8}\left\{\frac{-\left(\frac{c^2}{c_2^2}+\xi_2\right)}{\left(\frac{a_1-G}{k}-\frac{G}{2}\right)}+1\right\}^2\left(\frac{a_1-G}{k}-\frac{G}{2}\right)\right\}\left\{\left(\frac{c^2}{c_2^2}+\xi_2\right)-\frac{1}{2}\left(\frac{a_1-G}{k}-\frac{G}{2}\right)-1\right\}+\frac{1}{8}\left\{\frac{-\left(\frac{c^2}{c_2^2}+\xi_2\right)}{\left(\frac{a_1-G}{k}-\frac{G}{2}\right)}+1\right\}^2\left(\frac{a_1-G}{k}-\frac{G}{2}\right)^2\right]}$$
(28)

where $\xi_2 = \frac{P_2}{2\mu_2}$

Equation (28) gives the dispersion equation of Love wave in the anisotropic porous medium of finite thickness H under traction free boundary and overlying elastic gravitating heterogeneous half-space under initial stress.

PARTICULAR CASE

When $\alpha = 1$ i.e. $L \rightarrow \mu_1, d \rightarrow 1, a_1 \rightarrow 0$ and $G \rightarrow 0$, upper layer is homogeneous and lower half-space is homogeneous without gravitational force the equation (27) becomes

$$\tan\left(\sqrt{\left(\frac{c^2}{\beta_1^2}-1\right)}kH\right) = \frac{\mu_2}{\mu_1} \frac{\left[\sqrt{1-\frac{c^2}{c_2^2}}\right]}{\left[\sqrt{\left(\frac{c^2}{\beta_1^2}-1\right)}\right]}$$
(29)

This is the dispersion equation of Love waves in a homogeneous isotropic half-space in the absence of rigid layer.

ANGE OF LOVE WAVE SPEED

From the equation (22), it follows that Love waves can propagate in the porous layer under initial stress overlying heterogeneous elastic half-space under gravity if

$$\beta_1\sqrt{\frac{1-\xi_1}{d}} < c < c_2$$
(30)

Relation indicates the roll of initial stress and porosity of the media for the existence and non-existence of Love waves.

NUMERICAL RESULTS AND DISCUSSION

The dispersion curves of Love wave is drawn by taking $d = 0.6, \frac{c}{c_2} = 0.87, \frac{L}{\mu_2} = 2.5$ and other data from Table 1 to Table 5.

In Figure 1 and Figure 2, the phase velocity of the Love wave in absence of gravity is more than in the presence of gravity in the presence of compressive and tensile initial stress in the upper layer in case of rigid boundary. In case of traction free boundary, the phase velocity of the Love wave is same under the initial stress of the upper layer in the presence or absence of gravity.

The phase velocity of the Love waves in absence of gravity are more than in the presence of gravity without initial stress of the upper layer in rigid boundary, but in case of traction free boundary, the phase velocity of Love wave is same in the presence or absence of gravity without initial stress.

Also we observed that the phase velocity of Love waves in rigid boundary is more than traction free boundary, either upper layer is presence or absence of gravity or initial stress.

In Figure 3, the phase velocity of Love wave in absence of gravity is more than in presence of gravity in presence of heterogeneity parameter in case of rigid boundary where as in traction free boundary, the phase velocity of Love wave is same.

If the heterogeneity parameter increases, the phase velocity of Love wave increases in presence or absence of gravity under rigid where as in traction free boundary the phase velocity decreases.

In Figure 4 and Figure 5, the phase velocity of Love wave in absence of gravity parameter is more than the presence of gravity parameter under compressive and tensile stress in half-space, under rigid boundary, whereas the phase velocity of Love wave is same in case of traction free boundary.

The phase velocity of Love wave in absence of compressive and tensile stress in the half-space is more than the presence of gravity parameter under rigid boundary where as in traction free boundary, the phase velocity of Love wave is same. The phase velocity of Love wave in rigid boundary is always more than the traction free boundary in presence or absence of compressive and tensile stress in the half space.

Table 1. Parameters of Figure 1

Rigid Boundary Curve No.	Traction free Boundary Curve No	ξ_1	G
1	5	0.3	0.3
2	6	0.0	0.3
3	7	0.3	0.0
4	8	0.0	0.0

Table 2. Parameters of Figure 2

Rigid Boundary Curve No.	Traction free Boundary Curve No.	ξ_1	G
1	5	0.0	0.3
2	6	-0.3	0.3
3	7	0.0	0.0
4	8	-0.3	0.0

Table 3. Parameters of Figure 3

Rigid Boundary Curve No.	Traction free Boundary Curve No.	G	$\frac{a_1}{k}$
1	5	0.3	0.01
2	6	0.3	0.03
3	7	0.0	0.01
4	8	0.0	0.03

Table 4. Parameters of Figure 4

Rigid Boundary Curve No.	Traction free Boundary Curve No.	ξ_2	G
1	5	0.0	0.3
2	6	-0.3	0.3
3	7	0.0	0.0
4	8	-0.3	0.0

Table 5. Parameters of Figure 5

Rigid Boundary Curve No.	Traction free Boundary Curve No.	ξ_2	G
1	5	0.0	0.3
2	6	0.3	0.3
3	7	0.0	0.0
4	8	0.3	0.0

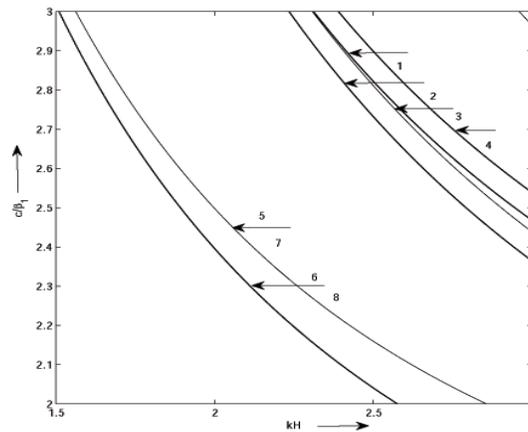


Figure 3. Variation of dimensionless phase velocity $\frac{c}{\beta_1}$ against the wave number kH for the different values of initial stress

ξ_1 .

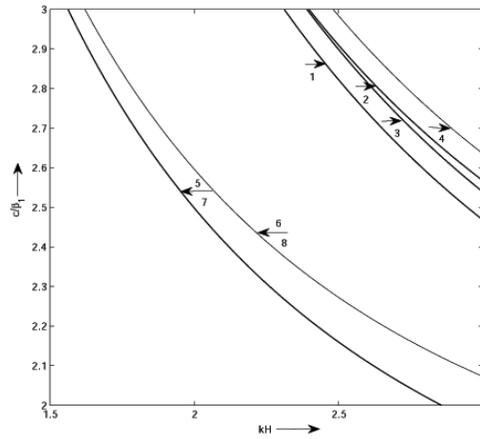


Figure 4. Variation of dimensionless phase velocity $\frac{c}{\beta}$ against the wave number kH for the different values of heterogeneity parameter.

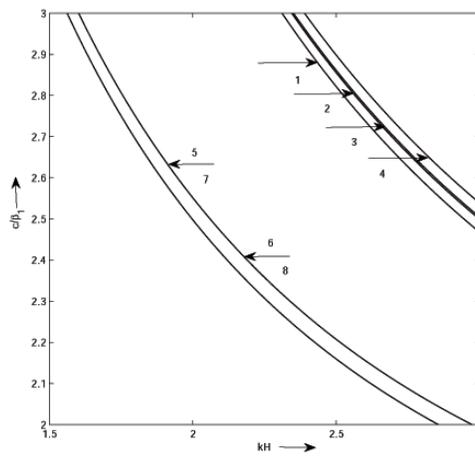


Figure 5. Variation of dimensionless phase velocity $\frac{c}{\beta}$ against the wave number kH for the different values of heterogeneity parameter and G .

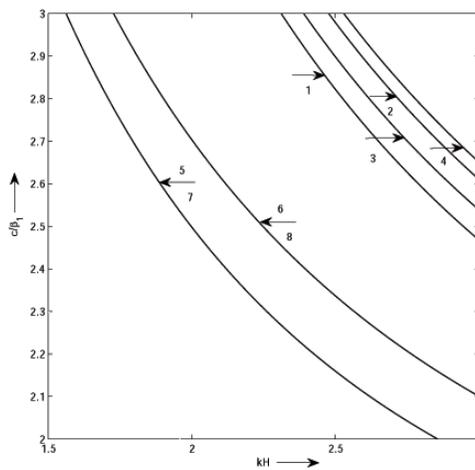


Figure 6. Variation of dimensionless phase velocity $\frac{c}{\beta}$ against the wave number kH for the different values of compressive initial stress s_2 and G .

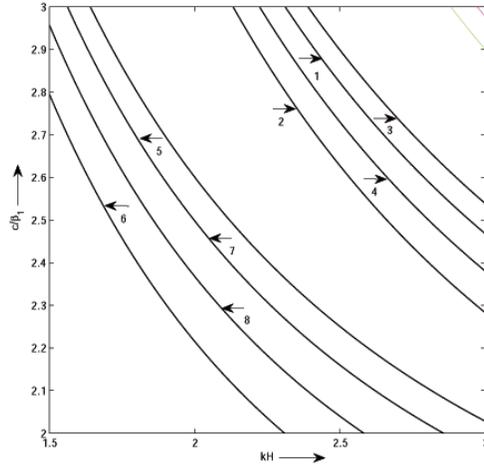


Figure 7. Variation of dimensionless phase velocity $\frac{c_p}{c_s}$ against the wave number kh for the different values of initial stress ϵ_2 and G_1 .

CONCLUSIONS

Propagation of Love waves in porous layer under initially stress over heterogeneous half-space under gravity and initially stress has been investigated analytically in rigid boundary as well as traction free boundary. The dispersion equation is obtained in both the cases. From the figures we may conclude that

1. In absence or presence of initial stresses in the half-space or layer, the phase velocity of Love waves in rigid boundary is more than the traction free boundary in presence or absence of gravity.
2. In presence of initial stress in the half-space and absence of initial stress in the layer, the phase velocities of Love wave in traction free boundary is more than the rigid boundary.
3. Initial stresses play a vital role on the propagation of Love waves in the assumed medium.
4. Gravity and heterogeneity in the half-space play a notable effect on the propagation of Love waves in the medium and half-space.

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CONFLICT OF INTEREST

The authors are stated that there is no conflict of interest between them.

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APPENDIX-I

Derivation from equation (2)-(3):

$$\begin{aligned} \left(N - \frac{P_1}{2}\right) \frac{\partial^2 v}{\partial x^2} + L \frac{\partial^2 v}{\partial z^2} &= \frac{\partial^2}{\partial t^2} [\rho_{11} v + \rho_{12} V] \quad \text{and} \\ \frac{\partial^2}{\partial t^2} [\rho_{12} v + \rho_{22} V] &= 0 \end{aligned} \quad (2)$$

Let $d_2 = [\rho_{12} v + \rho_{22} V]$, then $V = (d_2 - \rho_{12} v) / \rho_{22}$

$$\text{Hence } d_1 = \rho_{11} - \frac{\rho_{12}^2}{\rho_{22}}$$

Therefore,

$$\left(N - \frac{P_1}{2}\right) \frac{\partial^2 v}{\partial x^2} + L \frac{\partial^2 v}{\partial z^2} = d_1 \frac{\partial^2 v}{\partial t^2}$$

(3)

Where, $\frac{\partial^2}{\partial t^2} \left[\frac{\rho_{12} d_2}{\rho_{22}} \right] = 0$, since $\frac{\rho_{12} d_2}{\rho_{22}}$ is a constant term.

APPENDIX –II

Derivation of equation (10):

we have

$$\frac{\partial S_{12}}{\partial x} + \frac{\partial S_{22}}{\partial y} + \frac{\partial S_{23}}{\partial z} - \rho g \omega_{23} + \rho g z \frac{\partial \omega_{12}}{\partial x} - \rho g z \frac{\partial \omega_{23}}{\partial z} - \frac{P_2}{2} \frac{\partial v}{\partial x^2} = \rho \frac{\partial^2 v_2}{\partial t^2} \quad (8)$$

$$S_{12} = 2\mu e_{xy}, \quad S_{23} = 2\mu e_{yz}, \quad e_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right), \quad e_{xy} = \frac{1}{2} \left(\frac{\partial v_1}{\partial y} + \frac{\partial v_2}{\partial x} \right), \quad e_{yz} = \frac{1}{2} \left(\frac{\partial v_2}{\partial z} + \frac{\partial v_3}{\partial y} \right)$$

In this problem $\frac{\partial}{\partial y} = 0$ and $\mu = \mu_2 (1 + a_1 z)$, $\rho = \rho_2$, the stress strain relation becomes

$$\begin{aligned}
 e_{xy} &= \frac{1}{2} \frac{\partial v_2}{\partial x}, \quad e_{yz} = \frac{1}{2} \frac{\partial v_2}{\partial z} \\
 S_{12} &= \mu_2 (1 + a_1 z) \frac{\partial v_2}{\partial x}, \quad S_{23} = \mu_2 (1 + a_1 z) \frac{\partial v_2}{\partial z}, \quad \frac{\partial S_{12}}{\partial x} = \mu_2 (1 + a_1 z) \frac{\partial^2 v_2}{\partial x^2} \\
 \frac{\partial S_{23}}{\partial x} &= \mu_2 (1 + a_1 z) \frac{\partial^2 v_2}{\partial z^2} + a_1 \mu_2 \frac{\partial v_2}{\partial z}, \quad \frac{\partial w_{12}}{\partial x} = -\frac{1}{2} \frac{\partial^2 v_2}{\partial x^2}, \quad \frac{\partial w_{12}}{\partial x} = \frac{1}{2} \frac{\partial^2 v_2}{\partial z^2}
 \end{aligned} \tag{9}$$

where, μ is the modulus of rigidity, a_1 is a variation parameter of a rigidity having dimension inverse of length.

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Equation (8) using the above relations takes the form

$$\left[\mu_2 (1 + a_1 z) - \frac{1}{2} \rho_2 g z - \frac{P_2}{2} \right] \frac{\partial^2 v_2}{\partial x^2} + \left[\mu_2 (1 + a_1 z) - \frac{1}{2} \rho_2 g z \right] \frac{\partial^2 v_2}{\partial z^2} + \left(a_1 \mu_2 - \frac{1}{2} \rho_2 g \right) \frac{\partial v_2}{\partial z} = \rho_2 \frac{\partial^2 v_2}{\partial t^2} \tag{10}$$