

Distribution of ($mb - Ms$) of moderate earthquakes along the Mexican seismic zone

Vyacheslav M. Zobin

Observatorio Vulcanológico, Universidad de Colima, 28045 Colima, Col., México.

Received: August 26, 1996; accepted: August 3, 1998.

RESUMEN

Se estudia la distribución espacial de los valores $mb-Ms$ para 55 sismos en el periodo de 1978 a 1994, con profundidades focales de 0 a 80 km, mb de 4.5 a 5.5, ocurridos a lo largo de la Trinchera Mesoamericana entre 95° y 107° W. Los valores $mb-Ms$ tienen una distribución bimodal, con picos en 0.3 y 0.8 y mínimo en $mb-Ms=0.6$. Los eventos del primer grupo, con $mb-Ms<0.6$, se distribuyeron a lo largo de la trinchera en toda la región, comprendida por los bloques Michoacán, Guerrero y Oaxaca; la mayoría de los eventos del segundo grupo, con $mb-Ms\geq 0.6$, se encontraron en la parte central de la región en el bloque de Guerrero. Estos datos muestran que la litosfera del bloque de Oaxaca está relativamente homogénea y genera los eventos de baja frecuencia, pero la litosfera del bloque de Guerrero está más compleja y produce los eventos de alta y baja frecuencia.

PALABRAS CLAVE: Temblor, magnitud, México, subducción.

ABSTRACT

We study the spatial distribution of $mb-Ms$ for 55 earthquakes from 1978 to 1994 with focal depth from 0 to 80 km, mb from 4.5 to 5.5, which occurred along the Middle America trench, between 95° to 107° W. The $mb-Ms$ values have a bimodal distribution, with peaks at values of 0.3 and 0.8 and minimum at $mb-Ms=0.6$. Events of the first group, $mb-Ms < 0.6$, were distributed along the trench over the whole region, and events of the second group, $mb-Ms \geq 0.6$, were found mainly in the central part of the region, within the Guerrero block. These data show that the Oaxaca block consists of a relatively homogeneous lithosphere, where low-frequency events dominate, while the Guerrero block is more complex radiating both low- and high-frequency events.

KEY WORDS: Earthquake, magnitude, Mexico, subduction.

INTRODUCTION

Magnitudes M_s and mb characterize the intensity of earthquake radiation in different frequency ranges. Magnitude M_s is based on the surface wave amplitude at a period around 20 s, while magnitude mb is determined by measuring teleseismic P-wave amplitude at a period of about 1 s. The difference between the mb and M_s values, $mb-M_s$, defines a relationship between short- and long-period earthquake radiation and may be an indicator of the tectonic environment of the seismic zone. This difference is widely used, for example, for discrimination between earthquakes and nuclear explosions. Normally, $mb-M_s$ is about 2 to 3 for explosions and 0 to 2, for earthquakes (Denny *et al.* 1987). Prozorov *et al.* (1983) and Prozorov and Sabina (1984) used M_s-mb values for the world-wide regionalization of the active seismic zones. They showed the possibility of distinguishing between subduction zones (low-frequency events), and ridges and fracture zones (high-frequency events) using the M_s-mb discriminator. Ekstrom and Dziewonski (1988) found a systematic difference in the M_s versus seismic moment (M_0) relationships for continental, and ridge and fracture zone earthquakes. They demonstrated the predominance of short-period radiation for continental earthquakes.

We study the $mb-M_s$ distribution along the Mexican seismic zone. This region is characterized by a complex tectonic setting. Transverse fracture zones and ridges cut the Middle American Trench (Figure 1). There are active blocks along the Mexican coast (Jalisco, Michoacán, Guerrero, Oaxaca) whose boundaries reflect these fracture zone-trench intersections. Each of these blocks is characterized by high seismic activity related to subduction of the young Rivera (Jalisco block) and Cocos (Michoacán, Guerrero, Oaxaca, blocks) plates beneath the North American plate (Pardo and Suárez, 1995).

The region of this study extends from 15° to 19° N and 95° to 107° W (Figure 2). We study only the events east of the Middle America trench, where the oceanic Cocos plate subducts under the North America plate. LeFevre and McNally (1985) noted that the region north of the Tehuantepec Ridge may be considered as homogeneously stressed.

OBSERVATIONS

Figure 3 shows a plot of magnitude mb versus M_s for 126 earthquakes from the Mexican seismic zone, focal depth

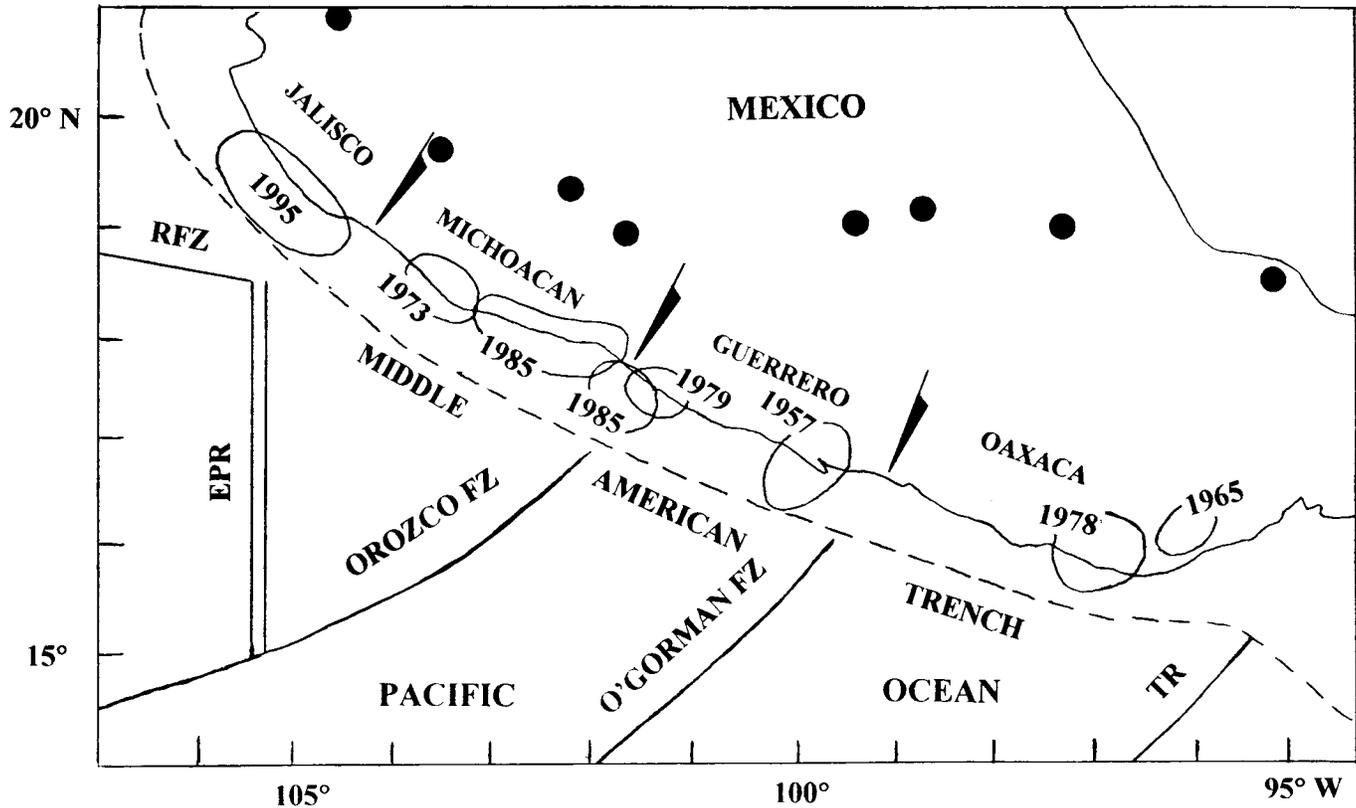


Fig. 1. Main tectonic structures of the region of study. Jalisco, Michoacán, Guerrero and Oaxaca, active blocks of the subduction zone after Singh and Mortera (1991). Black circles indicate the position of active volcanos. A dashed line shows the axis of the Middle America trench. Ovals show the rupture zones of the great earthquakes, $M_s \geq 7.5$ for the last 50 years and the year of event, after Singh and Mortera (1991) and Zobin (1997). The double line shows the East Pacific Rise (EPR), the single lines show the main oceanic Fracture Zones (FZ) and Tehuantepec Ridge (TR). Arrows show the boundaries between active blocks. RFZ, the Rivera Fracture Zone.

0 to 80 km, which occurred from 1962 to 1993. Magnitudes m_b and M_s were taken from the Bulletin of the International Seismological Centre (1964-1993) and Monthly Listings, NEIC (1962,1963). Data are plotted for events with magnitude m_b less than 6.5.

A maximum-likelihood linear regression for $4.5 \leq m_b \leq 6.5$ and $3.0 \leq M_s \leq 6.5$ yields

$$M_s = 2.49 (\pm 0.05) m_b - 7.96 (\pm 0.27) \quad R = 0.85 \quad (1)$$

Here R is the coefficient of correlation. The coefficients of the equation are close to those obtained by Prozorov and Sabina (1984) for Mexico using data from 1964 to 1980.

According to this equation, $m_b \sim 5.5$ may be considered as a boundary value; for lower values of m_b , $m_b \geq M_s$, while for higher values, $M_s > m_b$. Geller (1976) suggested that m_b may saturate beyond 5.5. To assure a homogeneous data set and to avoid the problem of magnitude saturation, we only consider events with $m_b \leq 5.5$.

We selected 55 events which occurred along the Middle America trench between 1978 and 1994, with focal depth

from 0 km to 80 km (Table 1). Epicenters, depth of events, and magnitudes m_b and M_s were taken from ISC (Bulletin of the International Seismological Centre, 1978-1993) and UNAM (Boletín Sismológico Mensual, Instituto de Geofísica, UNAM, 1994). A map of epicenters is presented in Figure 2. Mean errors in the event locations, taken from ISC (93% of events), are 6.3 km for epicenter and 6.6 km for depth. Errors in the determination of events by UNAM (4 of 55 events) were not published. Standard deviation in the determination of M_s is 0.08 for a group of stations and 0.22 for a single station; standard deviation in the determination of m_b is 0.06 for a group of stations and 0.34 for a single station (Kaverina *et al.*, 1996). Most of the magnitudes were determined from several stations. The magnitudes m_b were estimated from data from 10 to 83 stations and magnitudes M_s were estimated from data from 1 to 10 stations (26 estimations were made by 3 or more stations). Prozorov and Hudson (1974) showed a slight dependence of $m_b - M_s$ upon focal depth. Our data support this result. The coefficient of correlation is 0.164 for the 55 events (not significant at the 90% confidence level). Therefore, we neglect the depth difference as a factor on $m_b - M_s$ for the depth interval of 0 to 80 km.

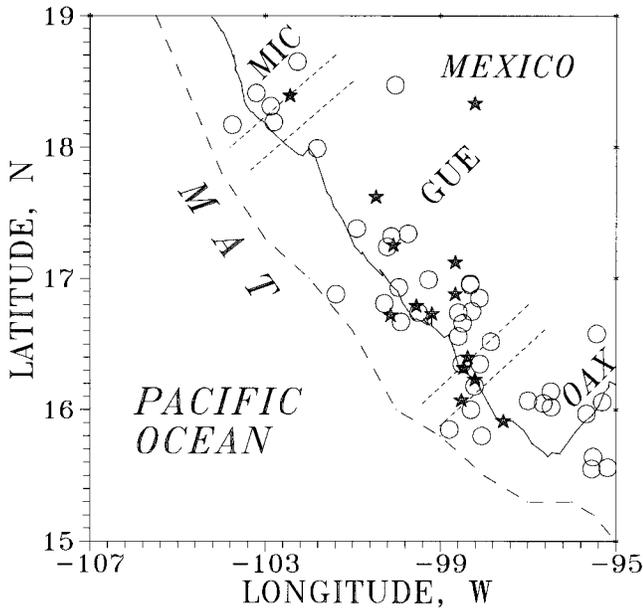


Fig. 2. Spatial distribution of events with different $mb-M_s$. Circles, $mb-M_s < 0.6$; stars, $mb-M_s \geq 0.6$. The dashed line shows the position of the axis of the trench (MAT). The parallel dashed lines show the borders between active blocks. MIC, the Michoacán, GUE, the Guerrero, OAX, the Oaxaca blocks.

Small events with magnitude $mb < 5.5$ appear to reflect the properties of the source better than the larger events because large events are generally complex (Singh and Mortera, 1991). Ekstrom and Dziewonski (1988) suggest that for large earthquakes the differences between short- and long-period radiation for different tectonic provinces are not as significant as for the smaller ones.

We find no dependence of $mb-M_s$ upon magnitude mb ($R = 0.13$), but there is a slight negative dependence upon M_s ($R = -0.75$). Therefore, small values of $mb-M_s$ indicate relatively large amount of low-frequency radiation (low-frequency events); while large values indicate small amount of low-frequency radiation by the earthquake source (high-frequency events).

DISTRIBUTION OF $mb-M_s$ ALONG THE MIDDLE AMERICA TRENCH

Figure 4 shows the percentage distribution of $mb-M_s$ for the 55 events, which suggests a bimodal distribution with peaks at $mb-M_s = 0.3$ and 0.8 . The significance of these two peaks was tested by Student's criterion for small samples (Caulcott, 1973):

$$t = \frac{(x_1 - x_2) \sqrt{[(n_1 - 1)s_1^2 + (n_2 - 1)s_2^2] (1/n_1 + 1/n_2)}}{[(n_1 - 1) + (n_2 - 1)]} \quad (2)$$

where t is the critical value of the criterion; x_1 and x_2 are the

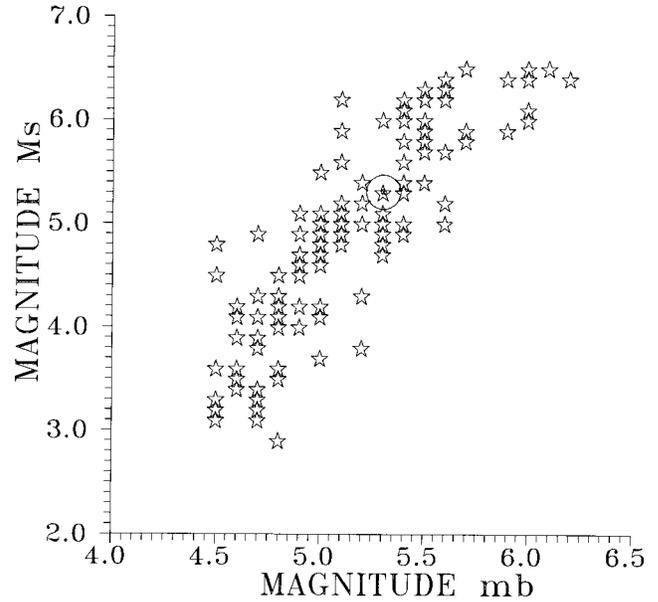


Fig. 3. Magnitude mb versus magnitude M_s plot for Mexican earthquakes, 1962-1993, depth range 0-80 km. The large circle shows the point where $mb = M_s$.

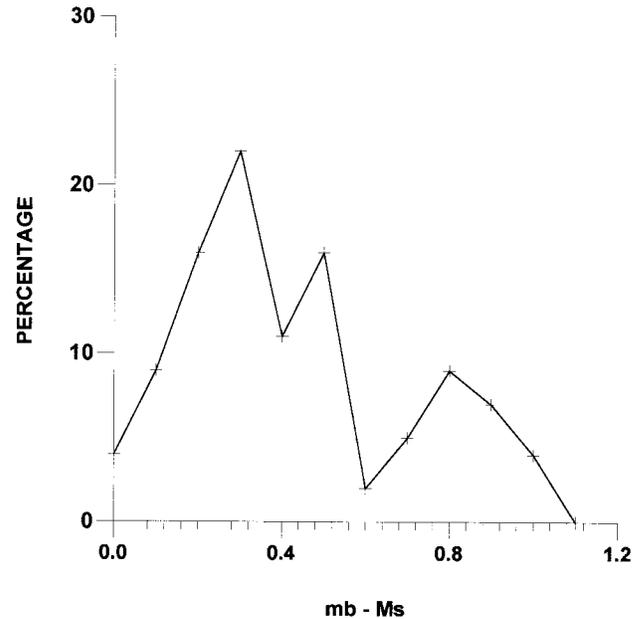


Fig. 4. The percentage distribution of the $mb-M_s$ values.

means of two samples, with magnitude difference from 0 to 0.5 and more than 0.5; n_1 and n_2 are numbers of events in the two samples, and s_1 and s_2 are the sample standard deviations. The test was done for two cases: for all magnitudes, and for magnitudes estimated by 3 or more stations. In both cases the values of the peaks were the same (Table 2) and were significant at the 90% confidence level.

Table 1

List of events

yy	m	d	Lat., N	Long., W	Dep., km	mb	Ms	<i>mb-Ms</i>
1978	07	05	18.47	-100.03	75	5.4	4.9	0.5
1978	09	29	18.65	-102.27	86	5.4	5.0	0.4
1978	11	30	16.07	-096.98	40	5.1	4.9	0.2
1978	12	11	16.67	-099.90	42	5.4	5.3	0.1
1978	12	28	16.02	-096.47	45	5.3	4.8	0.5
1979	01	06	18.32	-102.88	56	5.1	4.8	0.3
1979	01	28	16.85	-098.11	56	5.1	4.8	0.3
1981	06	20	16.07	-098.52	22	5.1	4.2	0.9
1981	07	21	16.40	-098.38	46	5.1	4.4	0.7
1981	07	21	16.35	-098.51	36	5.2	5.0	0.2
1982	01	02	16.81	-100.28	26	5.4	5.1	0.3
1982	01	02	16.72	-100.14	19	5.3	4.5	0.8
1982	12	14	16.66	-098.49	40	5.1	5.0	0.1
1983	02	07	16.96	-098.32	60	5.5	5.0	0.5
1983	02	21	16.95	-098.32	53	5.3	4.8	0.5
1984	01	28	17.32	-100.12	58	5.0	4.7	0.3
1984	06	04	18.33	-098.21	53	5.4	4.3	0.9
1984	07	14	17.34	-099.74	48	5.3	4.8	0.5
1984	10	19	16.75	-098.28	46	5.1	4.9	0.2
1984	11	30	16.23	-098.21	36	5.3	4.5	0.8
1984	12	13	16.14	-096.48	45	5.3	4.9	0.4
1985	09	25	18.19	-102.81	30	5.3	5.2	0.1
1986	05	29	17.12	-098.66	53	5.2	4.2	1.0
1987	06	07	16.88	-098.66	41	5.5	4.7	0.8
1987	09	27	18.17	-103.75	15	5.0	4.8	0.2
1988	09	01	16.73	-099.19	52	5.0	4.1	0.9
1989	05	02	16.99	-099.27	52	5.3	4.9	0.4
1989	10	08	17.25	-100.07	58	5.0	4.1	0.9
1989	11	09	16.79	-099.55	10	5.1	4.1	1.0
1990	01	13	16.74	-099.50	10	5.2	4.8	0.4
1990	01	29	18.39	-102.44	60	5.3	4.5	0.8
1990	02	17	15.80	-098.05	10	5.1	4.9	0.2
1991	01	14	17.99	-101.82	55	5.3	4.9	0.4
1991	04	01	16.35	-098.10	40	5.1	4.8	0.3
1991	04	27	17.24	-100.21	59	4.6	4.1	0.5
1991	05	22	18.41	-103.21	45	5.0	4.6	0.4
1991	07	25	16.88	-101.39	22	5.4	5.4	0.0
1991	09	10	16.06	-095.31	48	4.8	4.7	0.1
1991	09	26	16.18	-098.22	25	5.2	4.7	0.5
1991	10	22	16.58	-095.44	46	4.6	4.3	0.3
1991	11	10	15.55	-095.55	20	4.8	4.6	0.2
1991	11	24	16.52	-097.85	44	5.2	5.1	0.1
1992	01	08	15.56	-095.19	4	5.0	4.2	0.2
1992	05	15	16.32	-098.47	8	4.8	4.2	0.6
1992	05	19	16.00	-098.30	25	4.7	4.2	0.5
1992	06	07	16.74	-098.59	51	5.3	4.9	0.4
1992	06	07	16.56	-098.59	42	5.2	4.9	0.3
1992	11	02	16.05	-096.65	46	4.8	4.5	0.3
1992	11	10	16.93	-099.95	30	4.5	4.2	0.3
1993	03	31	17.38	-100.91	30	5.3	5.1	0.2
1993	07	29	17.62	-100.47	66	5.0	4.3	0.7
1994	07	09	15.91	-097.56	21	4.2	3.4	0.8
1994	08	06	15.64	-095.52	11	5.0	4.7	0.3
1994	09	17	15.97	-095.67	15	4.8	4.6	0.2
1994	12	13	15.85	-098.79	16	5.3	4.9	0.4

Note. 1978-1993, data from ISC; 1994, data from SSN, UNAM.

Table 2

Statistical test of significance for two peaks in the *mb-Ms* distribution

Test	n_1	x_1	s_1	n_2	x_2	s_2
1	41	0.304	0.139	14	0.828	0.113
2	26	0.308	0.138	7	0.800	0.115

t -critical at 90% confidence level of significance for 53 and 31 degrees of freedom is 1.68 and 1.70, respectively; t_1 is 13.10, t_2 is 8.63. t_1 and t_2 are $\gg t$ -critical.

We chose the minimum value between the two peaks in *mb-Ms* distribution, $mb-Ms = 0.6$, as a boundary between two groups: 14 high-frequency events ($mb-Ms \geq 0.6$) and 41 low-frequency events ($mb-Ms < 0.6$).

A map of epicenters (Figure 2) shows some regularity in space distribution for these two groups of events. Low-frequency events are situated along the trench of the whole Pacific coast of Mexico while the high-frequency events cluster mainly in the middle part of the region in the Guerrero block. To test this observation we constructed a 2x2 contingency table (Caulcott, 1973) for events situated within the Guerrero and Oaxaca blocks (Table 3). The 6 events occurring on the boundary between the two blocks (Figure 2) were excluded from this calculation. The value of χ^2 criteria was obtained from a contingency table (Caulcott, 1973):

$$\chi^2 = (p_1 - p_2)^2 / p(1-p)(1/n_1 + 1/n_2), \quad (3)$$

where $p_1 = x_1/n_1$; $p_2 = x_2/n_2$; $p = (x_1 + x_2)/(n_1 + n_2)$; n_1 and n_2 are the total number of events in the blocks of Oaxaca and Guerrero, respectively; and x_1 and x_2 are the number of events with $mb-Ms < 0.6$ in these blocks. This value of χ^2 was compared with χ^2 at the 90% confidence level. The test shows (Table 3) that there is a significant difference in occurrence of earthquakes with different spectral content in the two blocks. The number of events observed in the Michoacán block was not sufficient for testing.

DISCUSSION

The *mb-Ms* distribution along the Mexican seismic zone shows a difference for the blocks of Guerrero and Oaxaca. This study was based on ISC locations. Singh and Lermo (1985) noted a tendency of systematic mislocation of the 1973-1982 ISC epicenters as compared to their locations determined by the local networks installed in the near-source region. It was observed that the epicenters were shifted by about 35 km towards NE, or almost perpendicular to the trench. This one-directional systematic shift perpendicular

Table 3

Statistical test of significance for change in the *mb-Ms* distribution in the Guerrero and Oaxaca zones (a 2x2 contingency table).

Zone	Oaxaca	Guerrero	Total
<i>mb-Ms</i> < 0.6	13	20	33
<i>mb-Ms</i> ≥ 0.6	1	9	10
Total	14	29	43

χ^2 -critical at 90% confidence level of significance is 2.70; our value of $\chi^2 = 3.30$, or $\chi^2 > \chi^2$ -critical.

to epicentral zone does not change our results. We can say that the Oaxaca block consists of a relatively homogeneous lithosphere, where low-frequency events dominate, while the Guerrero block is more complex, radiating low-frequency as well as high-frequency events.

The change in physical properties at the boundary between the Guerrero and Oaxaca blocks is supported by other seismological investigations. Singh and Mortera (1991) report that large earthquakes are relatively simple in Oaxaca and complex in Guerrero. They also note that the ratio of surface-wave to body-wave seismic moments for large earthquakes is smaller in Oaxaca than in the regions northwest of 99° W. Singh and Mortera (1991) propose that the existence of this boundary may be related with the different age of the subducting plate which older in Oaxaca (about 20 m.y.) and younger in Guerrero (about 13 m.y.). Castro *et al.* (1994) observe a difference in the quality factor Q for Oaxaca and Guerrero blocks which indicates greater attenuation in Oaxaca than in Guerrero.

CONCLUSIONS

The *mb-Ms* values of small earthquakes along the Middle America trench, between 95° and 107° W, have a significant bimodal distribution, with peaks at 0.3 and 0.8. Events of the first group, *mb-Ms* < 0.6, were distributed along the trench in the whole region, but events of the second group, *mb-Ms* ≥ 0.6, were clustered mainly in the central part of the region within the Guerrero block. This suggests that the Oaxaca block consists of a relatively homogeneous lithosphere, where low-frequency events dominate, while the Guerrero block is more complex radiating both low- and high-frequency events.

ACKNOWLEDGMENTS

I am grateful to R. Zúñiga for providing me with a catalogue of the Mexican earthquakes of 1899-1990 and E.

Glowacka for providing with programs for a work with catalogue. Careful comments by R. Castro, A. Nava and C. Rebollar and two anonymous reviewers helped to improve the text. S.K. Singh helped in revision of the manuscript. This research was partly supported by the project 641132 of CICESE.

BIBLIOGRAPHY

- CASTRO, R. R., L. MUNGUÍA, C. J. REBOLLAR and J. G. ACOSTA, 1994. A comparative analysis of the quality factor Q for the regions of Guerrero and Oaxaca, Mexico. *Geofis. Int.*, 33, 373-383.
- CAULCOTT, E., 1973. Significance Tests. Routledge and Kegan, London and Boston, 128 p.
- DENNY, M. D., S. R. TAYLOR and E. S. VERGINO, 1987. Investigation of mb and Ms formulas for the western United States and their impact on the Ms/mb discriminant. *Bull. Seismol. Soc. Amer.*, 77, 987-995.
- EKSTROM, G. and A. M. DZIEWONSKI, 1988. Evidence of bias in estimations of earthquake size. *Nature*, 332, 319-323.
- GELLER, R., 1976. Scaling relations for earthquake source parameters and magnitudes. *Bull. Seismol. Soc. Amer.*, 66, 1501-1523.
- KAVERINA, A. N., A. V. LANDER and A. G. PROZOROV, 1996. Global creepex distribution and its relation to earthquake source geometry and tectonic origin. Report to IUGG Gener. Ass., Boulder, Col., USA.
- LEFEVRE, L. V. and K. C. McNALLY, 1985. Stress distribution and subduction of aseismic ridges in the Middle America subduction zone. *J. Geophys. Res.*, 90, 4495-4510.
- PARDO, M. and G. SUAREZ, 1995. Shape of the subducted Rivera and Cocos plates in southern Mexico: seismic and tectonic implications. *J. Geophys. Res.*, 100, 12357-12373.
- PROZOROV, A. G. and J. A. HUDSON, 1974. A study of the magnitude difference Ms-mb for earthquakes. *Geophys. J. R. A. S.*, 39, 551-564.
- PROZOROV, A. G., J. A. HUDSON and M. SHIMSHONI, 1983. The behaviour of earthquake magnitudes in space and time. *Geophys. J. R. A. S.*, 73, 1-16.

- PROZOROV, A. G. and F. J. SABINA, 1984. Study of the properties of seismicity in the Mexico region. *Geophys. J. R.A.S.*, 76, 317-336.
- SINGH, S.K. and J. LERMO, 1985. Mislocation of Mexican earthquakes as reported in international bulletins. *Geoffs. Int.*, 24, 333-351.
- SINGH, S. K. and F. MORTERA, 1991. Source time functions of large Mexican subduction earthquakes, morphology of the Benioff zone, age of the plate, and their tectonic implications. *J. Geophys. Res.*, 96, 21487-21502.
- SIGH, S. K., M. RORDRIGUEZ and L. ESTEVA, 1983. Statistics of small earthquakes and frequency of occurrence of large earthquakes along the Mexican subduction zone. *Bull. Seismol. Soc. Amer.*, 73, 1779-1776.
- ZOBIN, V. M., 1997. The rupture history of the Mw 8.0 Jalisco, Mexico, earthquake of 1995 October 9. *Geophys. J. Int.*, 130, 220-228.

Vyacheslav M. Zobin
*Observatorio Vulcanológico,
Universidad de Colima,
Colima, Col., 28045, México.
E-mail: vzobin@cigic.ucol.mx*