

Seismicity of the State of Puebla, Mexico, 1986-1989

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Received: June 22, 1994; accepted: February 15, 1995.

RESUMEN

En este trabajo se presenta un análisis de la actividad sísmica del estado de Puebla y sus inmediaciones, registrada durante el periodo de enero 1986 a agosto de 1989 por la Red Sísmica del Estado de Puebla (RESEP), perteneciente a la Universidad Autónoma de Puebla (UAP). La localización epicentral de los eventos fue determinada mediante la aplicación de los programas HYPO71PC y HYPOCENT. Se calcularon los errores teóricos en la localización de eventos mediante el programa HYPOERR los cuales coinciden con los errores estimados en las localizaciones. El análisis de los 350 eventos localizados con un residual cuadrático menor de 2, muestra que la mayoría de eventos tienen una profundidad menor de 50 km con excepción de la región sur de la zona de estudio, en la cual ocurren también sismos con profundidades mayores de 50 km. Estos últimos parecen estar asociados al contacto entre la litosfera oceánica y la continental que corresponden a las placas de Cocos y Norteamericana, respectivamente. El 92% de los eventos estudiados tiene una magnitud de coda entre 2 y 4, aunque es importante resaltar que en el área de estudio se han registrado eventos con magnitud mayor de 7.0.

PALABRAS CLAVE: Sismicidad, hipocentro, magnitud, estado de Puebla.

ABSTRACT

The seismic activity in the State of Puebla, Mexico and its surroundings, between January 1986 and August 1989, is discussed using data from the seismic network of the University of Puebla. The hypocentral locations were obtained using the programs HYPO71PC and HYPOCENT. The calculated theoretical errors using the program HYPOERR coincide with the estimated errors from the location procedure. A total of 350 earthquakes were located with rms less to 2. Most of the hypocenters are at depths less than 50 km, with the exception of the southern region, where events deeper than 50 km are found. These deeper events are probably associated with the contact between the oceanic and continental lithospheres, which correspond to the Cocos and North American plates. 92% of the events have coda magnitude between 2 and 4, although events with magnitudes greater than 7 have occurred in this region.

KEY WORDS: Seismicity, hypocenter, magnitud, Puebla State.

INTRODUCTION

Studies of the seismic activity in a region contribute to the determination of seismogenic zones and the elaboration of seismic zonation maps for engineering and disaster management purposes (Gutiérrez *et al*, 1991). The State of Puebla is located in a region with intermediate depth seismicity, as in the Orizaba earthquake of 1973 ($M=7.3$) and the October 24, 1980 Huajuapán de León earthquake ($M=7.0$), (González-Ruiz, J., 1986). The State is also located relatively close (≈ 350 km) to other zones of major seismic activity in south-east Mexico (Figure 1). The State of Puebla has four major urban centers, including the state capital, with a strong population, economic and industrial growth.

The State of Puebla is located between 17.8° and 20.8° N, and between 96.0° and 99.7° W. To the north and to the east, it shares boundaries with the State of Veracruz. The Sierra Norte de Puebla is a continuation of the Sierra Madre Oriental, the Cofre de Perote mountain, the Derrumbadas mountain zone and the Citlaltépetl volcano. To the south, the State borders with the States of Guerrero and

Oaxaca where we find the Sierra Negra, the Mixteca Poblana and the continuation of the Sierra Madre Occidental in the Sierra Madre del Sur, its most important feature is the Tehuacán fault. To the west it borders with the States of Morelos and Mexico, where the Popocatepetl and Iztaccihuatl strato-volcanoes are located, and with the States of Tlaxcala and Hidalgo where the Malintzi volcano is situated (Figure 2). The State is traversed by five lower Miocene fault systems (Mooser, 1972); the more important ones are Popocatepetl-Chignahuapan, Malintzi and Atoyac-Minas. The seismic activity in the region may be due to tectonic movements or to volcanism.

An early seismic study of the State of Puebla and its surroundings is by Figueroa (1974). His catalog covered from 1523 to October of 1973, including historic investigations of major earthquakes up to 1910, plus instrumental earthquakes with magnitudes greater than 3 from 1911 to October 1973. He reports a total of 231 earthquakes; 108 historical events and 123 instrumental earthquakes.

Gómez and González-Pomposo (1983), report 104 earthquakes of all magnitudes from 1976 to February 1983.

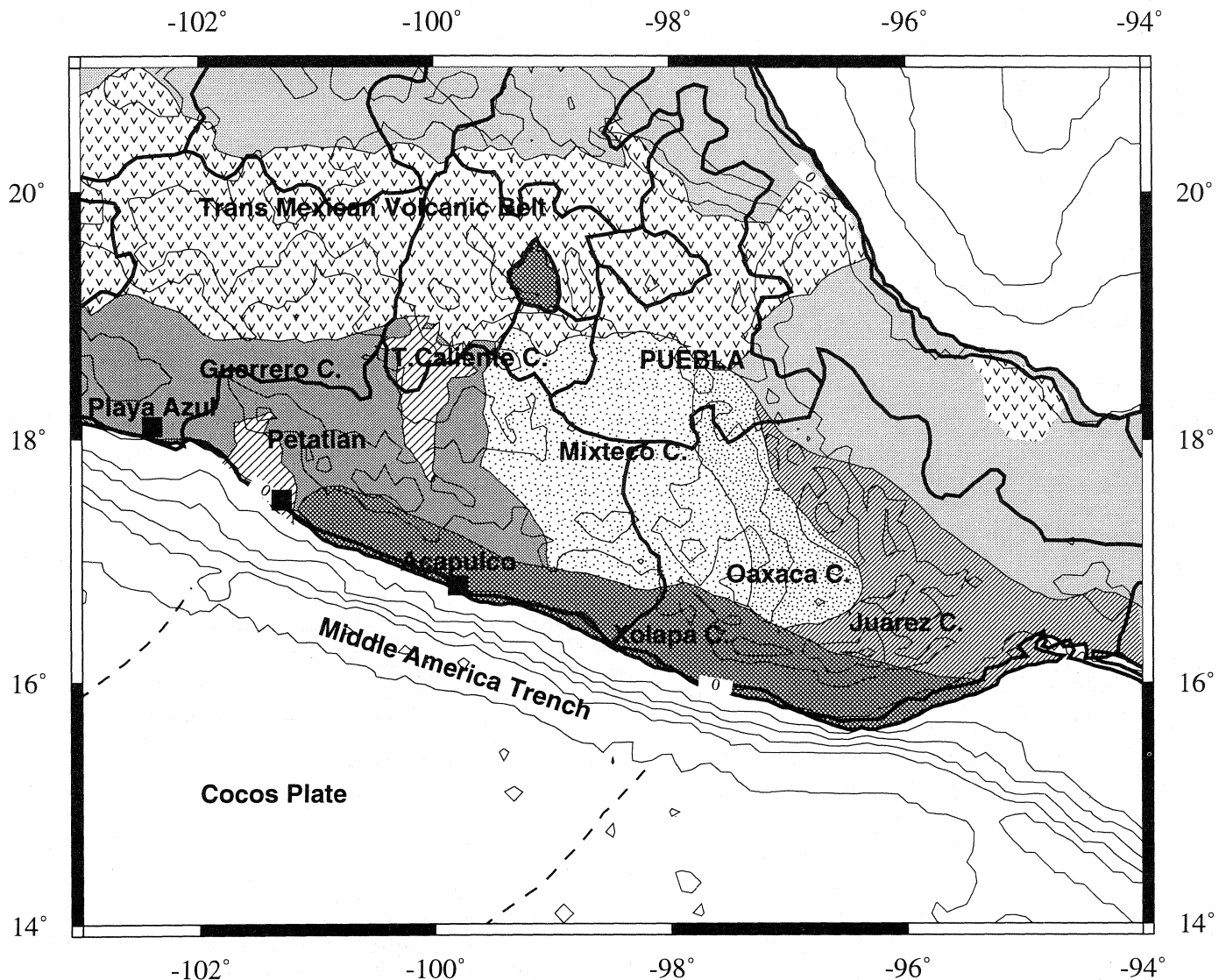


Fig. 1. Location map of the study area.

Thus all previous studies report only 227 earthquakes between 1911 and February 1983. This small number may be attributed mainly to the lack of seismic instrumentation in the region.

On November 29, 1985 the Seismic Network of the State of Puebla (RESEP) began operation. Initially RESEP had three seismic stations, in Chiautla de Tapia (CHP), Molcaxac (MOP) and Cuyoaco (CUP) (Serrano and González-Pomposo, 1985). In 1986 a station was added in Coxcatlán (CXP). In 1987 RESEP installed three more stations in Chila de las Flores (CIP), Ciudad Universitaria (UAP) and Xalitzintla (XLP) (González-Pomposo *et al.*, 1987). These seven seismic stations operated through

August 1989 (Figure 3). The geographical coordinates of the RESEP stations, and of supporting stations belonging to SISMEX (IIT, IIS, IIA, IIC, III), plus those belonging to the National Seismic Network, (VHO, OXM, LVM and TPM) are shown in Table 1. All together, these stations provide a wide coverage and adequate control of hypocentral determinations in the region of interest.

INSTRUMENTATION

The equipment used in the RESEP stations was analogic, one-component Vertical and with smoke paper recording. Each station had a Ranger SS-1 seismometer

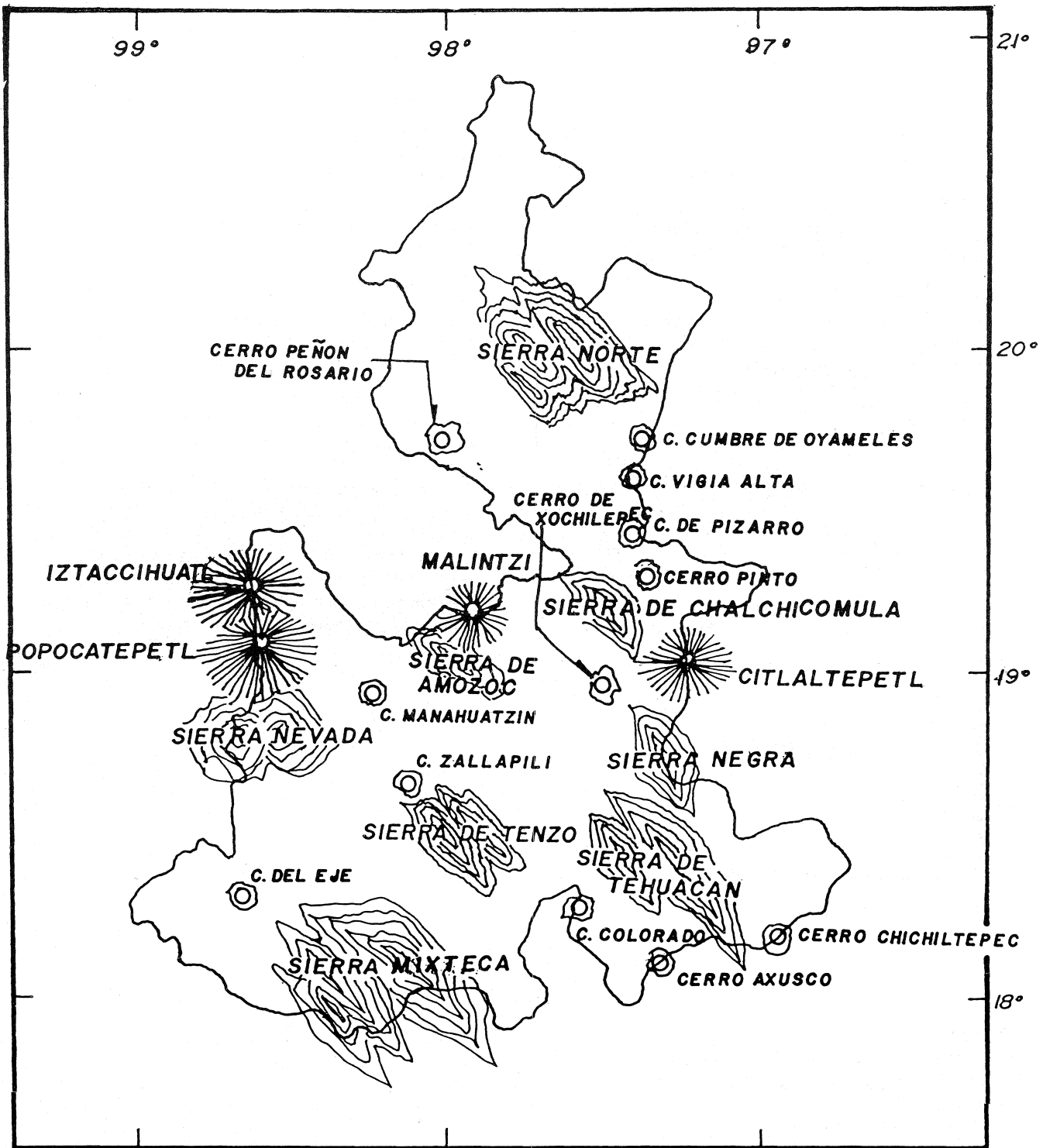


Fig. 2. Orography of State of Puebla.

with a natural period of 1 sec, with an MEQ-800-B seismograph, and a UTC radio time receiver in the 5, 10 and 15 MHZ frequency band. The housing for the equipment was designed at the School of Civil Engineering, and the power systems were built at the Department of Electronic Maintenance of the School of Physics and Mathematics,

also of the University of Puebla. The stations CHP, CUP, CIP, and XLP, are on bedrock, while for stations MOP, UAP, and CXP are on a one cubic-meter concrete piers due to poorly consolidated ground. The seismic stations were visited once a week for recollection of data and for maintenance up to 1989 (Serrano and González-Pomposo, 1985).

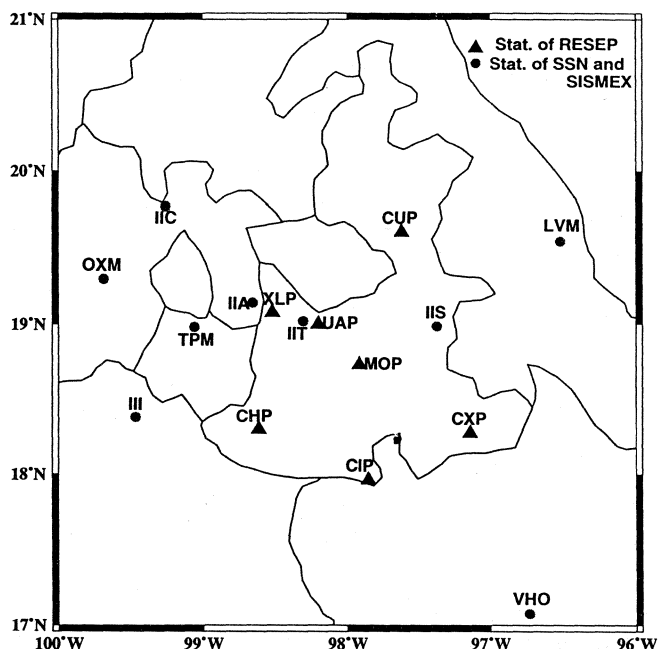


Fig. 3. Stations of RESEP.

Table 1

Coordinates of stations of RESEP				
NAME	Code	Lat. N	Log. W	Altitude (m)
Chiautla de Tapia, Pue.	CHP	18.298	98.613	1030
Chila de las Flores, Puebla.	CIP	17.963	97.854	1765
Cuyoaco, Puebla.	CUP	19.603	97.619	2450
Coxcatlán, Puebla.	CXP	18.274	97.142	1280
Molcaxac, Puebla.	MOP	18.733	97.917	1840
Ciudad Universitaria, Pue.	UAP	19.001	98.203	2114
Xalitzintla, Puebla.	XLP	19.079	98.521	2600

Coordinates of station of SSN and SIXMEX				
NAME	Code	Lat. N	Log. W	Altitude (m)
Alzomoni, Edo. de México.	IIA	19.143	98.655	3900
Santa Rita Coyotepec	IIC	19.757	99.258	2725
Iguala, Guerrero	III	18.376	99.468	1750
Ciudad Serdán, Puebla.	IIS	18.986	97.372	
Tonanzintla, Puebla.	IIT	19.021	98.308	2205
Laguna Verde, Veracruz	LVM	19.613	96.395	160
Oxotlán, Edo. de México.	OXM	19.296	99.688	2700
Tepoztlán, Morelos.	TPM	18.983	99.061	1500
Observatorio, Oaxaca Oax.	VHO	17.079	96.732	1685

DATA PROCESSING

Locations of event were evaluated with the program HYPOERR (Lienert *et al.*, 1986). This program provides uncertainties estimated for a specific station distribution with a given layered model. It calculates vertical and horizontal uncertainties in hypocentral locations, as well as origin-time uncertainties. An analysis was carried on with

all 16 stations in Table 1 to determine the precision in origin time, latitude, longitude and depth determination for events located within the region 96.2 to 99.2°W and 17.3° to 21.0°N, and 0 to 60 km depth at 10 km intervals.

The data collected by RESEP were processed as follows. First, the arrival times for the P and S waves (T_p and T_s) were read as well as they characteristics, the signal duration, and maximum amplitude of each record. We applied the criteria that the seismic events should have a $\Delta t = T_s - T_p$ of less than 20 seconds at three stations for the event to be included as belonging to the region. Records from SISMEX and SSN stations were included. If the s-p difference was larger, the readings were sent to SSN to be used in the location of regional events.

The events were then located using the computer location program HYPO71 (Lee and Stewart, 1975). This program calculates in an iterative way the hypocentral parameters. The program uses an analytic expression of first arrival travel times in a flat layered model using non-linear least squares. For the hypocentral determinations we also used HYPO-PC (Lee and Valdés, 1985). This program was tested with more than 20 stations and with a seismic model of up to 10 flat layers. In order to improve the hypocentral determination, we also used HYPOCENTER (Lienert *et al.*, 1986) with the seismic model in Table 2.

Table 2

Layered velocity model	
Thickness (km)	Velocity (km/s)
0 - 5	5.00
5 - 20	6.10
20 - 25	6.95
25 - 45	7.60
45 - 00	8.10

The flat layer model used in the present study (Table 2) is in agreement with the one proposed by Valdés *et al.* (1986). This model was constructed by P-wave ray-tracing in two dimensions. The seismic profile was oriented north-south across the central part of the region of study. The explosions were located in Lake Alchichica, north of the city of Puebla, and off the Coast of the Pacific Ocean. This model provided hypocentral determinations with the smallest residuals. In comparison with the layer models from SISMEX, SSN and RESMAC the hypocentral locations were less scattered and had lower residuals. Layered seismic models based on other seismic profiles (Nuñez-Cornú, 1988; Singh and Pardo, 1993) were also tested, but found to be inadequate.

Coda magnitudes were computed after Havskov and Macías (1982)

$$M_c = 0.09 + 1.85 \log(Z) + 0.0004(D),$$

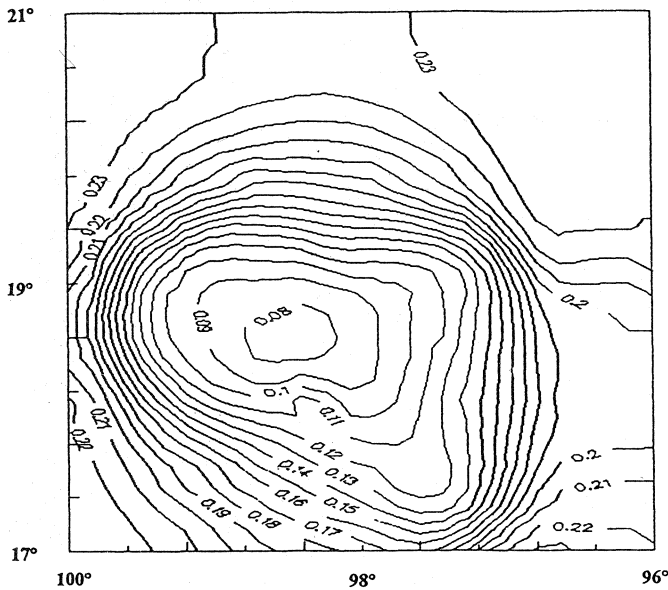


Fig. 4. Origin time (error in sec.).

where Z is the signal duration in seconds and D is the epicentral distance in km.

RESULTS

The results for the evaluation of the seismic network using the program HYPOERR at a depth of 30 km, shows that for the central region ($\sim 98.5^\circ$ to 99.5° W and 18.5° to 19.5° N) the origin time error is less than 0.11 seconds. For the outer area the error is up to 0.23 seconds (Figure 4). The uncertainties in the latitude and longitude determination at a depth of 30 km and for the same regions values of 1.4 km and 3.5 km were calculated, respectively. In the case of the depth uncertainty, the values obtained were 1.4 km and 1.8 km for the same regions as before (Figure 5).

The maximum errors found in hypocentral determination for the central region was 1 to 3 km error in the depth calculation, and 9 to 11 km error for the outer region. The origin time errors were in the range of 0.3 to 0.4 seconds. Larger estimated errors in the north region are due to the sparse network coverage. The south region has the highest seismic activity. The evaluation of the uncertainty in the hypocenter parameter determination indicates an acceptable error range.

Figure 6 shows the seismic events located with the program HYPOCENT, as well as their errors in the latitude and longitude estimation. These errors disagree with those calculated by HYPOERR, because not all 16 stations used in HYPOERR recorded P and S waves arrivals for all events. There were a total of 580 of events registered by RESEP between January 1986 and August 1989. From these, only 350 (60 %) have an RMS error smaller than 2

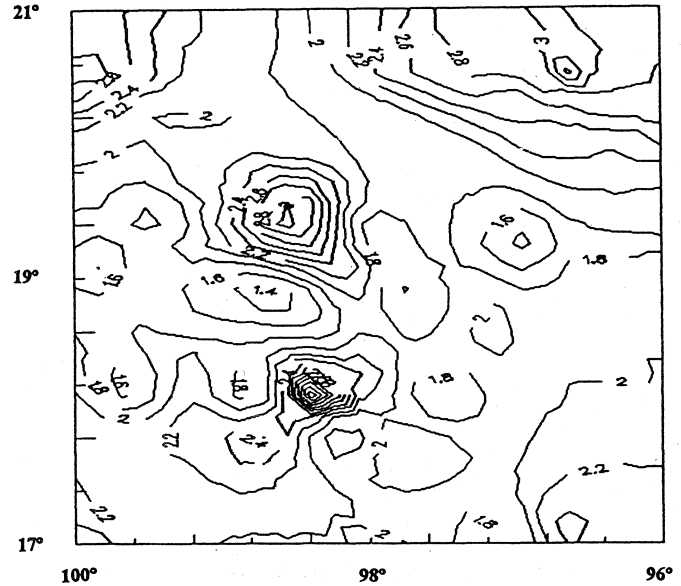


Fig. 5. Depth (error in km.).

seconds in origin time. Table 3 shows the hypocentral parameters.

Among 350 events, 43% had magnitudes between 2 and 3, 49% between 3 and 4, 8% between 4 and 5, and only one event greater than 5. This event is on the edge of the region of study. Figure 7 shows the magnitude vs number of events distribution. The catalog appears to be complete for magnitudes between 1.5 and 3.3. The b value estimated is 1, which is the expected value for shallow tectonic events (Bullen and Bolt, 1993).

Figure 8 shows a histogram of the monthly seismicity for the period of time studied. We observe an increment of seismicity starting in June, 1986 and reaching a peak in June, 1988, then a decrease until August 1989 when the network stopped operation. The months with high activity are: June, 1986 (17 events); January, 1988 (16 events), June, 1988 (28 events), September, 1988 (17 events), February 1989, (19 events) and March, 1989 (15 events).

Between 96° and 100° W and 17° and 18.5° N, we located 241 earthquakes (69%). This corresponds to the Chiautla, Tehuacán, Acatlan, border of Puebla-Guerrero and Puebla-Oaxaca active regions. Thus the seismicity of the State of Puebla is concentrated in the South region. In the Central zone (18.5° to 20° N) only 26% of the events are located; this includes the regions of Zacatlán, Atlixco, Puebla City, Serdán and the border between Puebla-Veracruz, a region of moderate seismic activity. At the borders of the States of Tlaxcala-Mexico-Hidalgo (98.3° to 99.1° W and 19.4° to 20° N), only 18 events were located, 11 during May, 1986. They are probably associated with the Tlaloc-Actopan fault system. In the region known as Derrumbadas, close to

Table 3 (Cont.)

no	Date	Hr: Min	Sec	LatN	Log W	Depth	Mag	rms
277	881210	2117	53.86	18.1983	97.6915	37	2.6	0.6
278	881212	643	40.30	17.2082	96.0435	18	3.8	1.2
279	881224	1732	54.79	16.7858	99.0988	32	4.3	1.8
280	881230	736	46.13	17.9567	98.3662	37	3.3	1.6
281	890101	36	1.47	17.9865	98.3553	37	2.2	0.2
282	890101	53	17.37	18.0917	98.0663	44	2.6	0.3
283	890101	93	4.67	18.0250	97.7878	41	3.6	0.5
284	890101	2332	28.67	17.5982	97.2697	50	3.1	0.1
285	890104	1320	6.03	17.9947	97.8422	49	3.0	0.0
286	890106	531	39.98	17.0547	98.9157	37	3.8	0.3
287	890114	2314	55.05	18.0698	97.9073	44	2.7	0.1
288	890115	17 8	33.87	17.9637	97.7228	37	2.4	0.8
289	890117	442	40.17	18.0858	98.2123	37	2.8	0.4
290	890118	4 4	32.84	17.9637	97.5710	40	2.8	0.7
291	890118	1643	19.46	17.5930	97.0255	34	3.0	0.4
292	890120	618	47.66	17.9657	97.3742	48	2.8	0.2
293	890122	530	19.02	18.0677	97.0527	37	3.2	1.5
294	890122	1633	15.55	17.7913	95.9080	41	4.0	1.3
295	890124	752	12.86	17.1785	96.9332	37	4.0	0.6
296	890124	1146	8.41	18.5705	99.5563	30	3.9	1.1
297	890125	05	2.20	16.2997	98.4878	37	4.1	0.9
298	890126	1555	8.40	17.3988	97.2583	37	3.2	0.3
299	890126	16 0	13.18	17.5822	97.0005	51	3.1	0.2
300	890201	1650	17.63	17.8792	97.8765	47	2.5	0.1
301	890201	2321	14.12	17.9637	97.6312	39	2.5	0.7
302	890203	454	20.65	17.8283	99.4697	37	2.5	1.2
303	890206	651	22.21	17.4395	97.2737	48	2.9	0.1
304	890207	1241	54.38	17.9453	97.9355	43	2.8	0.0
305	890209	14 6	59.31	18.0902	98.7662	40	3.5	0.8
306	890209	2149	53.70	17.9928	99.1067	33	2.3	0.4
307	890212	1120	9.85	17.9178	98.6942	37	3.1	0.6
308	890213	149	4.21	19.1678	95.9627	24	3.8	0.1
309	890224	20 0	56.21	19.0588	98.6148	37	2.4	1.5
310	890226	1413	8.42	18.0837	95.9550	41	3.5	0.6
311	890227	2011	25.43	17.9467	98.6837	43	2.7	0.3
312	890313	330	59.90	16.8142	99.6528	37	3.8	0.5
313	890313	16 0	16.99	16.2342	98.2430	37	4.0	1.9

Citlaltépetl volcano, 13 events were located as shown in Figure 6. North of the State no events were located. Figure 6 shows the overall seismicity between January, 1986 and August, 1989.

Figure 9 is a north-south profile of the seismicity in the region. The vertical and horizontal hypocentral errors are indicated, as well as the boundary of the oceanic and continental lithosphere, proposed by Valdés *et al.* (1986). The suggested Benioff zone from Bevis and Isacks (1984) and Pardo (1993) were also compared, but these two profiles either fell out of the region of study or did not agree with the hypocenters in the present study. Only 28 events occur beneath the Benioff zone. The deep seismicity ends at about 18.8°N. A concentration of events is observed at 18°N at about 45 km depth, which matches the location of the Benioff zone.

CONCLUSIONS

Ninety-two percent of the seismic events located in the region of the State of Puebla have coda magnitudes ranging between 2 and 4. Only one event, slightly outside of the region of study, had a magnitude greater than 5. This suggests that the earthquakes in this region are mainly small.

Tabla 3 (Cont.)

no	Date	Hr: Min	Sec	LatN	Log W	Depth	Mag	rms
314	890316	2121	58.09	17.6120	98.0462	41	3.8	2.0
315	890318	749	53.39	17.9637	98.0837	41	2.7	2.0
316	890320	2022	29.03	18.0345	97.7945	49	2.6	0.1
317	890321	97	34.66	17.8880	98.3725	37	2.2	0.2
318	890325	1750	15.35	18.1315	98.1425	37	2.8	0.6
319	890327	1717	37.29	18.1697	98.8403	37	2.6	0.3
320	890401	548	43.69	17.9895	97.8873	40	2.8	0.4
321	890402	225	27.14	19.0877	99.3035	18	2.7	0.6
322	890403	833	26.80	18.8875	98.7517	29	3.2	1.1
323	890403	1129	37.97	18.1233	99.2930	33	4.0	2.0
324	890404	1328	27.83	19.0227	98.7057	18	3.5	1.3
325	890404	1854	0.88	18.3522	98.0278	37	2.5	1.0
326	890408	18 4	54.07	16.9543	98.4625	29	4.0	1.3
327	890412	547	14.37	17.9637	97.5187	64	2.4	1.1
328	890412	1040	14.98	19.3355	97.4428	37	3.0	1.1
329	890412	1042	3.41	19.6053	97.2820	32	2.9	1.3
330	890420	741	25.16	18.9315	99.9492	29	2.6	0.2
331	890420	830	1.42	18.7477	97.5202	42	2.2	1.5
332	890425	2234	15.86	17.8127	99.3987	37	3.3	0.4
333	890428	929	39.44	16.7128	99.4498	37	3.2	0.2
334	890428	1641	59.78	18.3145	97.6100	41	2.3	1.7
335	890503	736	1.87	18.1068	97.7102	37	2.3	0.3
336	890504	011	29.80	17.2342	99.5718	32	3.2	0.2
337	890508	1114	39.04	17.8693	97.7782	47	2.5	0.2
338	890528	1712	58.48	17.2782	99.7218	37	3.2	0.2
339	890529	4 8	9.56	17.6435	98.9272	37	3.5	0.3
340	890630	1919	10.63	17.8143	97.2815	41	2.7	1.4
341	890707	2212	12.83	17.7297	97.5310	37	2.7	0.4
342	890708	627	45.09	19.3375	97.6995	32	3.0	1.2
343	890713	1354	5.94	18.0802	98.3837	37	3.2	1.7
344	890716	1159	18.34	19.1433	99.5712	37	3.2	1.5
345	890724	1030	1.63	17.9032	97.9605	58	3.4	1.9
346	890727	1237	14.36	17.7940	99.1638	37	3.4	1.2
347	890813	23 9	24.36	18.1077	97.0500	58	2.9	0.2
348	890816	329	30.34	16.9168	98.1487	37	3.2	0.0
349	890818	029	5.61	19.3533	97.2443	37	2.7	1.7
350	890828	2025	52.80	16.4120	98.0048	37	3.4	1.4

Some earthquakes in the region attain large magnitudes, such as the Orizaba earthquake ($M=7.0$) of 1973 and the Huajuapán de León earthquake ($M=7.0$) of 1980 (González-Ruiz, 1986). About 69 % of the seismicity is located south of 18.5°N, basically in the region of the Mixteca range near the boundary between the States of Puebla-Guerrero and Puebla-Oaxaca. In the central region, between 18.5° and 20°N, 26 % of the events was located, corresponding to the Trans-Mexican Volcanic Belt. Several clusters of earthquakes occurred at the boundary between the States of Tlaxcala, Mexico and Hidalgo, where 18 shallow events with maximum magnitudes of 4.2 were located plus a cluster of 13 events in the region of Derrumbadas, near Citlaltépetl volcano.

Deep seismicity occurs in the south of the State, probably related to the contact between the Cocos and North American plates as suggested by the seismic profile of Valdés *et al.* (1986). In the central region, seismicity is shallower than in the south. It does not appear to be related to this contact, but to crustal tectonic stresses.

The larger seismic risk in the State of Puebla may be due to the subduction zone in the Pacific coast. This risk is now being studied using records from the Accelerograph

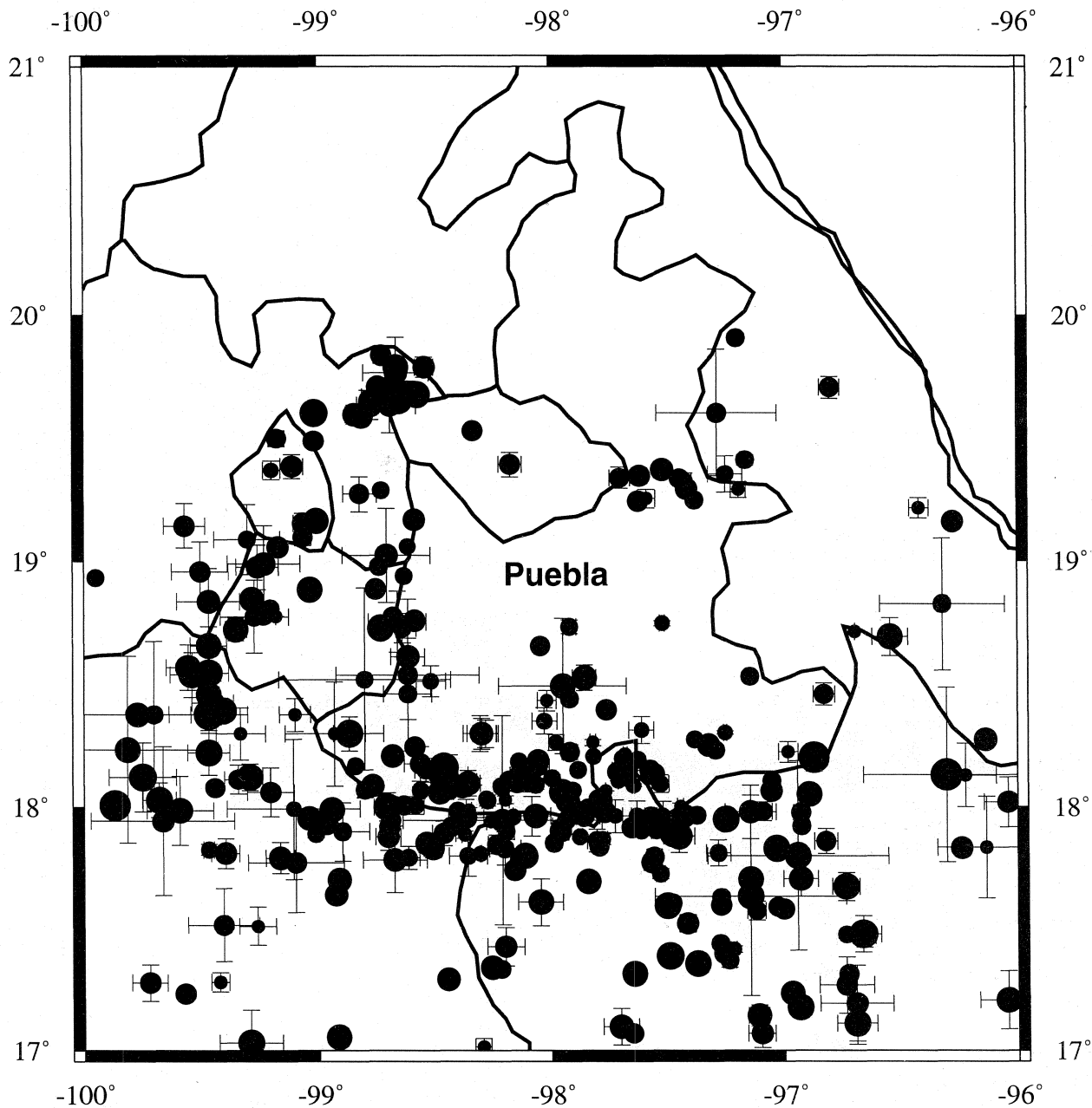


Fig. 6. Seismic epicentral distribution in the State of Puebla 1986-1989.

Network of Puebla City (Avilés *et al.* 1989). It is important to continue monitoring the seismicity in the State of Puebla, in order to understand the seismic hazard in the region and to contribute to the mitigation of destructive seismic effects.

ACKNOWLEDGEMENTS

We thank Dr. Raúl Serran L. and Eng. Jorge A. De Gante González for their contribution to the installation, maintenance and operation of RESEP; Patricia Gertrudis Juárez Serran, Franco Martínez Sánchez and Gustavo Adolfo Diego Taboada for their participation in reading the seismograms; the students of social service in the School

of Civil Engineering at the UAP who participated and contributed in many different ways to the operation of RESEP; the Instituto Tecnológico de Oaxaca for providing the station Coxcatlán in 1986; and Ana Elena Posada Sánchez for her assistance in the elaboration of the seismic catalog.

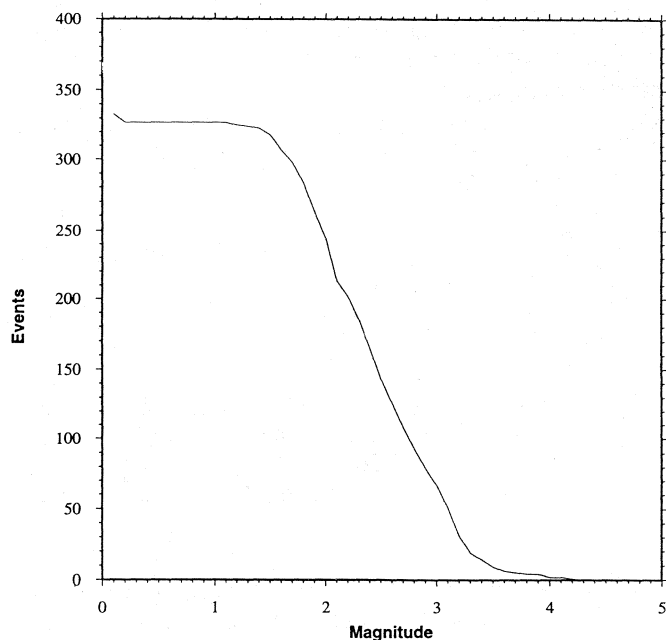


Fig. 7. Number of events versus magnitude of Coda.

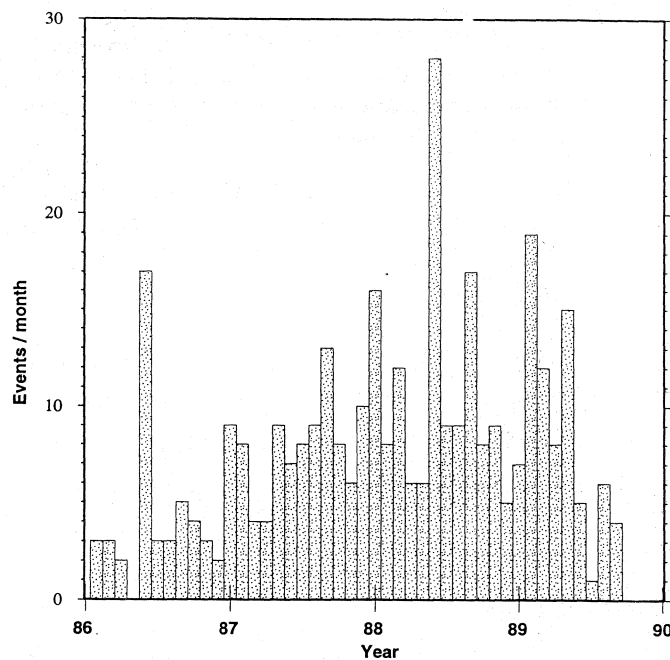


Fig. 8 Monthly seismicity histogram.

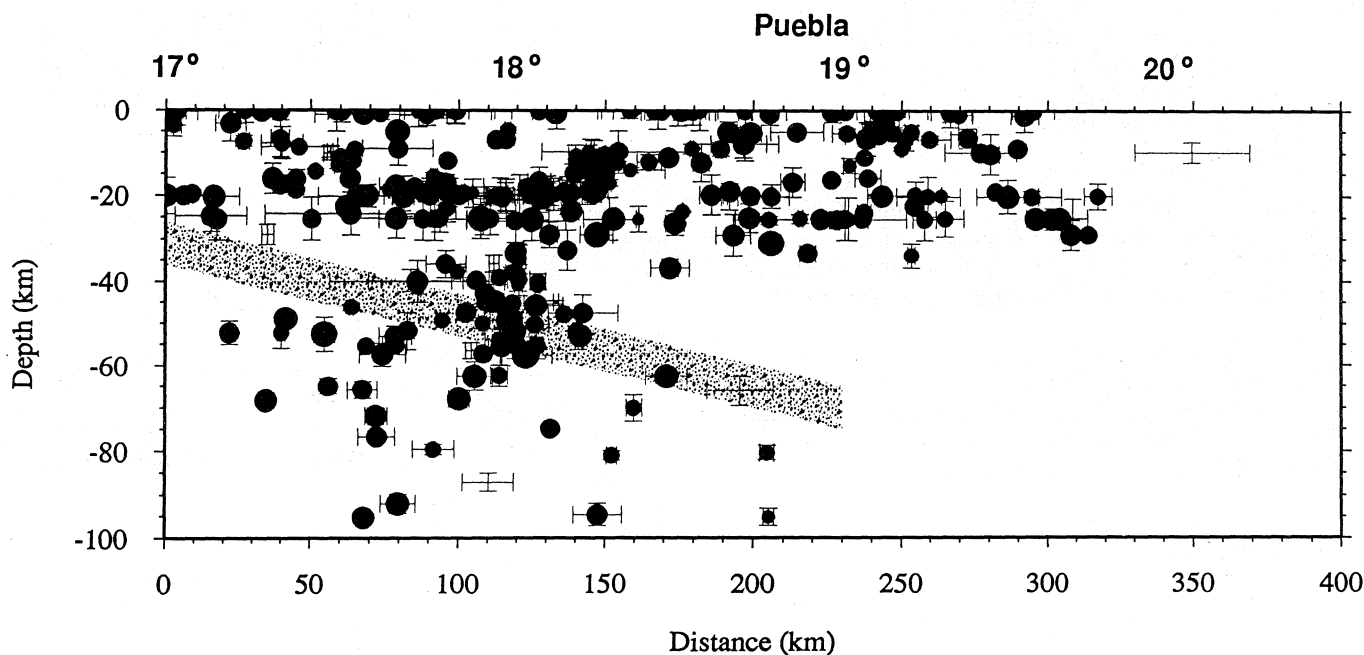


Fig. 9 Seismic cross section.

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